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ORIGINAL ARTICLE

Riverbed methanotrophy sustained by high carbon conversion efficiency

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Our understanding of the role of freshwaters in the global carbon cycle is being revised, but there is still a lack of data, especially for the cycling of methane, in rivers and streams. Unravelling the role of methanotrophy is key to determining the fate of methane in rivers. Here we focus on the carbon conversion efficiency (CCE) of methanotrophy, that is, how much organic carbon is produced per mole of CH₄ oxidised, and how this is influenced by variation in methanotroph communities. First, we show that the CCE of riverbed methanotrophs is consistently high (~50%) across a wide range of methane concentrations (~10–7000 nM) and despite a 10-fold span in the rate of methane oxidation. Then, we show that this high conversion efficiency is largely conserved (50% ± confidence interval 44–56%) across pronounced variation in the key functional gene (70 operational taxonomic units (OTUs)), particulate methane monooxygenase (*pmoA*), and marked shifts in the abundance of Type I and Type II methanotrophs in eight replicate chalk streams. These data may suggest a degree of functional redundancy within the variable methanotroph community inhabiting these streams and that some of the variation in *pmoA* may reflect a suite of enzymes of different methane affinities which enables such a large range of methane concentrations to be oxidised. The latter, coupled to their high CCE, enables the methanotrophs to sustain net production throughout the year, regardless of the marked temporal and spatial changes that occur in methane.

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Introduction

Freshwaters are now recognised as important components of the global carbon cycle (Cole *et al.*, 2007; Battin *et al.*, 2009) but there is a scarcity of data for streams and rivers (Bastviken *et al.*, 2011). Although small in their surface area, accounting for just 3–5% of inland water cover, some 40% of the 2.7 Pg of carbon conveyed to them each year never reaches the ocean; instead, it is mineralised to CO₂ and CH₄ (Battin *et al.*, 2009). Methane is the more potent of the two carbon greenhouse gases, and although efforts to quantify its outgassing from rivers are useful (Garnier *et al.*, 2013; Crawford *et al.*, 2014; Sawakuchi *et al.*, 2014), we know very little about its cycling in rivers. Proper integration of rivers into the global carbon cycle will not only require an understanding of the specific processes that determine their emissions of methane but also the microorganisms that mediate both its production and consumption (He *et al.*, 2012; Nazaries *et al.*, 2013).

The vast majority of methane produced in freshwater ecosystems, such as wetlands, lakes and peatbogs, does not escape to the atmosphere. Instead, it is oxidised to CO₂ (King *et al.*, 1990; Nedwell and Watson, 1995; Bastviken *et al.*, 2002) by predominantly aerobic methanotrophic bacteria, which use methane as their sole source of energy and carbon (Trotsenko *et al.*, 2008; He *et al.*, 2012). Methanotrophy, therefore, not only has the potential to markedly alter the balance of carbon gases (CO₂+CH₄) released from any aquatic ecosystem but it is also a form of chemosynthetic production (Trotsenko *et al.*, 2008). Such chemosynthetic production is now known to be important in many lakes across the globe (Jones and Grey, 2011) and recently we highlighted its widespread potential throughout many rivers across southern England, which are oversaturated in methane relative to the atmosphere (Shelley *et al.*, 2014).

Although the methane-oxidising bacteria responsible for such chemosynthetic production have been identified in peatlands (Chen *et al.*, 2008), lakes (Surakasi *et al.*, 2010; He *et al.*, 2012; Saidi-Mehrabad *et al.*, 2013) and wetlands (Yun *et al.*, 2012), they remain largely unknown in streams and rivers (Dumestre *et al.*, 2002; Chen and Murrell, 2010). To the best of our knowledge, there is no information on the respective efficiency of carbon

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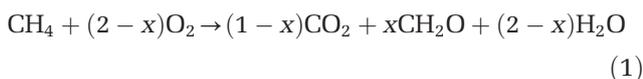
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conversion by any of these aquatic methanotroph communities. The latter will not only help to sustain the methanotrophic potential of any aquatic ecosystem to modulate emissions of methane but it will also govern the contribution of methane as a chemosynthetic energy source to the wider food web.

To date, most freshwater ecologists would argue that riverine food webs are based firmly on a combination of allochthonous detrital resources and autochthonous photosynthetic production, with only the balance between the two being contentious. However, our findings show that, even at methane concentrations in the low nanomolar range (typically 30–200 nM), riverbed methanotrophy can also contribute significant organic production (Trimmer *et al.*, 2010; Shelley *et al.*, 2014). Furthermore, substantial quantities (Gg C y^{-1}) of methane are known to be oxidised in large riverine systems, including the Amazon and the Hudson River (De Angelis and Scranton, 1993; Melack *et al.*, 2004).

In order to better understand methanotrophy in streams and rivers, we aimed to quantify its carbon conversion efficiency (CCE), that is, what fraction of methane oxidised is assimilated into biomass and whether that efficiency is influenced by variation in the methanotroph communities. During methanotrophy, the fraction of methane assimilated (x) can be represented as (Urmann *et al.*, 2009):



Although it is comparatively easy to measure an amount of methane oxidised by a sample of soil, water or sediments, actually converting this to an amount of carbon assimilated or fixed (x in equation (1)) has proved to be a non-trivial task (King *et al.*, 1990; Bastviken *et al.*, 2003; Maxfield *et al.*, 2012). Original work with $^{14}\text{C-CH}_4$ in high-methane, land-fill cover soils (Whalen *et al.*, 1990), reported that 69%

of methane oxidised was recovered in the bulk organic carbon fraction and, in low-methane, forest soils, a lesser fraction of 31–43% (Roslev *et al.*, 1997), whereas in the oxic water column of some lakes, the range is even more extreme (6–77%; Bastviken *et al.*, 2003). Such changes in the efficiency with which methanotrophs harness the energy in methane will govern their production and, ultimately, understanding the CCE of methanotrophy may help to explain changes seen in the ratio of CO_2 to CH_4 emitted from some wetlands and lakes (Bastviken *et al.*, 2003; Yvon-Durocher *et al.*, 2014).

The recent development of a stable isotope switching technique ($^{13}\text{CH}_4/^{12}\text{CH}_4$), in combination with sophisticated multi-compound-specific mass spectrometry, has begun to probe the true dynamics of carbon cycling in soils, but it cannot give a genuine measure of the efficiency of organic carbon production via methanotrophy (Maxfield *et al.*, 2012). Here we propose a simple alternative approach to quantifying the CCE of methanotrophy by quantifying the fraction of $^{13}\text{C-CH}_4$ recovered as $^{13}\text{C-DIC}$. We have already characterised a gradient in the potential for methanotrophy across 32 chalk streams in southern England (Shelley *et al.*, 2014); we now examine the CCE of methanotrophy across a subset of these streams and seek to determine whether that efficiency is conserved across any natural variation in the methanotroph community.

Materials and methods

Study sites and gravel collection

Eight typical English chalk streams were selected from a larger suite of survey sites (Shelley *et al.*, 2014) for which we already had the methane concentration, methane oxidation and water chemistry data (see Figure 1 main text and Supplementary Figure 1). We began by tracing $^{13}\text{C-CH}_4$ into $^{13}\text{C-organic}$ and

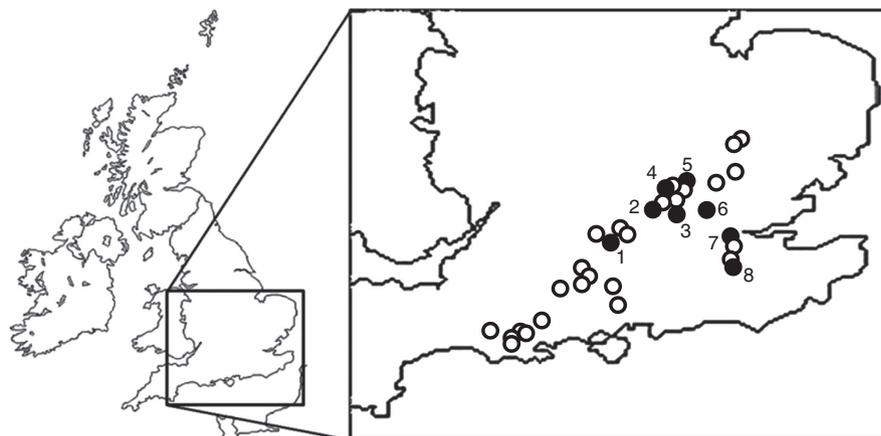


Figure 1 Distribution of the 32 (open and closed circles) chalk streams in southern England that have all proved positive for methane oxidation (see Shelley *et al.* (2014) for full details). Closed circles indicate the location of the eight streams used in this study, where the total distance east to west was approximately 110 km. 1. Lambourn, 2. Misbourne, 3. Chess, 4. Ver, 5. Gade, 6. Mimram, 7. Cray and 8. Darent.

^{13}C -inorganic carbon in repeat batch incubations using gravels from the River Mimram (51.80524 N–0.151212 W, Hertfordshire). Then, due to site access restrictions, for the follow-up set of ^{13}C -DIC yield experiments, we used gravels from the River Lambourn (51.438585 N–1.384889 W, Berkshire) where we had also worked previously (Trimmer *et al.*, 2009, 2010). Once the method for determining the CCE was optimised and validated, we broadened our work out to the full eight streams. Surface layer riverbed gravels ($n=5$) were sampled using a kick net and stored in a portable refrigerator for return to the laboratory where they were pooled and mixed, and then rinsed and sieved ($1.4\text{ mm} \leq x < 5.66\text{ mm}$) to remove any invertebrates (Trimmer *et al.*, 2009).

Tracing ^{13}C - CH_4 into ^{13}C -organic and ^{13}C -inorganic carbon in repeat batch incubations

Subsamples ($\sim 8\text{ g}$) of gravel were weighed into serum bottles (20 ml) along with air-equilibrated river water (10 ml) and then sealed, leaving an air headspace. The vials were enriched with ^{13}C - CH_4 (99 atom%) to give 445 p.p.m. in the headspace and 613 nM ($\pm 50\text{ nM}$, s.e., $n=8$) in the water phase. Higher concentrations for methane than ambient enable shorter incubations and have no effect on the measure of efficiency (see below). Additional control vials were set up with just river water. All vials were incubated on a tipper in the dark at 8°C (to mimic average river water temperature). The headspace was then measured repeatedly (every 8–12 h) for methane by GC/FID (Sanders *et al.*, 2007) and the change in methane over the first one or two time points was used to calculate the rate of methane oxidation normalised to the dry mass of gravel. Once $>90\%$ of the methane had been oxidised in each batch, the headspace was analysed for ^{13}C - CO_2 by continuous flow isotope ratio mass spectrometry (CF/IRMS), and then the water was removed, acidified ($50\ \mu\text{l HCl } 12.2\text{ M}$), and measured for ^{13}C as DIC (ΣCO_2 , HCO_3^- and CO_3^{2-}) by CF/IRMS (Miyajima *et al.*, 1995; Trimmer *et al.*, 2009). Summation of ^{13}C - CO_2 plus ^{13}C -DIC equalled the total amount of ^{13}C - CH_4 metabolised to inorganic carbon. After each incubation, a subsample of the gravel (2 g) was harvested, freeze dried and stored at -18°C for analysis of ^{13}C in either the lipid or bulk organic carbon (TOC) fraction (see below). New water was added to the remaining gravel and the vials enriched again with CH_4 . This was repeated six more times to give eight batches in total. Thus, the labelling of TOC on the gravel was a cumulative measure and, as both headspace and water were replaced, the DIC was discrete for each incubation.

Quantifying the yield of bulk ^{13}C -organic carbon by wet oxidation to CO_2

An adaptation of a standard wet oxidation method (no. 505C; (APHA, 1995)) was used to recover and quantify the yield of bulk ^{13}C -organic carbon from

the harvested gravels (see Supplementary Methods and Supplementary Figure 2 for recovery and sensitivity of the wet oxidation assay). After oxidation, the headspace was analysed for CO_2 , first by GC/FID (as above but after catalytic reduction to CH_4 over hot nickel), to quantify the TOC and then by CF/IRMS for ^{13}C - CO_2 to quantify the labelled fraction (as above).

Quantifying the yield of ^{13}C -lipids

A crude lipid extract was performed on the freeze-dried gravels (see Supplementary Methods for details) that would have included other hydrophobic compounds, such as lipopolysaccharides (Gurr and Harwood, 1991; Parrish *et al.*, 2000). After recovery of the organic layers, the extracts were pooled, concentrated (to $250\ \mu\text{l}$) under N_2 and stored at -18°C until later combustion and elemental analysis coupled to CF/IRMS (Trimmer *et al.*, 2009). Analytical controls of just UHP water were added to every batch.

Quantifying the CCE by the yield of ^{13}C -DIC (ΣCO_2 , HCO_3^- and CO_3^{2-})

After establishing that we could recover 100% of the ^{13}C - CH_4 as either ^{13}C -DIC or ^{13}C -organic carbon (bulk plus crude lipid), we could use the ^{13}C -DIC fraction to more simply quantify the efficiency of organic carbon production, without the need to quantitatively extract and purify the organic fractions (see Supplementary Table 1 and Results). If the production of ^{13}C -DIC is linear over time, with an origin through zero, then all of that ^{13}C -DIC must be solely due to methanotrophic metabolism. Consequently, the proportion of ^{13}C - CH_4 recovered as ^{13}C -DIC yields a direct measure of the CCE, for example, $1 - (\Delta^{13}\text{C-DIC}/\Delta^{13}\text{C-CH}_4) = x$ in equation (1) and Δ is the production or consumption of ^{13}C -DIC or ^{13}C - CH_4 , respectively. Here we take the total amount of ^{13}C - CH_4 oxidised to represent gross methanotrophy, whereas our measure of CCE is equivalent to net methanotrophy, that is, the amount of fixed carbon that would potentially be available to higher trophic levels and so is synonymous with net photosynthetic production (Shelley *et al.*, 2014).

Accordingly, we followed the oxidation of ^{13}C - CH_4 and parallel evolution of ^{13}C -DIC during short (60 h) incubations. Fifty-five aliquots of prepared gravels were enriched with 99 atom% ^{13}C - CH_4 to give 1900 p.p.m. in the headspace and 2400 nM in the water (± 65 , s.e., $n=55$), incubated as above and then sacrificed (five vials \sim every 6 h) for the quantification of CH_4 and ^{13}C -DIC (as above). Next, we tested the effect of methane concentration on the oxidation kinetics and conversion efficiency at 37 different concentrations (~ 10 – 7000 nM CH_4 in the water), spanning the seasonal range in the river water, gravel porewater (~ 30 – 150 nM) and far beyond (Trimmer *et al.*, 2009; Shelley *et al.*, 2014). The headspace was sampled repeatedly over time

(as above) and analysed for CH₄ and ¹³C-CO₂ (ΣDIC) to calculate the CCE (as above).

Finally, although the use of 99 atom% ¹³C-CH₄ assures the detectable labelling of products during short incubations (<5 h), fractionation can occur during methanotrophy (Whiticar, 1999). In order to test for any effect of fractionation, prepared gravels were incubated under 11 different mixtures of ¹³CH₄/¹²CH₄: from natural abundance (here simply ¹²CH₄) to 99 atom% ¹³CH₄ in ~10% increments, and at a final concentration in the water phase of 635 nM (±32 nM, s.e., *n* = 10).

Streambed gravels as a natural source of copper

Copper has a well-characterised (Semrau *et al.*, 2010) and strong overall influence on methanotrophs (physiology, gene expression, proteome) and, in particular, expression and activity of particulate methane monooxygenase (pMMO). We, therefore, tested whether the riverbed was a natural source of copper by incubating gravels with river water (as above), gravel with UHP water, river water only and UHP only, at 8 °C, as above, for 4 days. The water phase was then filtered (prerinsed 0.2 µm Mini-Sart) and analysed for copper by furnace atomic absorption spectrometry (Varian GTA 110 & 220FS, Victoria, Australia) calibrated between 0 and 21 µg Cu l⁻¹.

CCE across a natural gradient of variation in methanotroph communities

Having established a simple routine assay for estimating the CCE of methanotrophy, we now wanted to estimate that efficiency using a larger set of replicated samples drawn from eight independent streams (Figure 1) and examine whether the efficiency was consistent across any natural variation in the methanotroph community. The CCE of methane oxidation for each sample was quantified as above. DNA was extracted from one sample from each stream using the physical lysis based method of McKew *et al.* (2007). Here the functional gene encoding particulate methane monooxygenase (*pmoA*) was amplified using PCR and variation in that *pmoA* was used as a proxy for variation in the methanotroph communities across our eight streams. It is important to appreciate that we were not trying to compare diversity or determine whether the active methanotroph community varied between the streams; instead, we were testing whether or not the CCE of methane oxidation was conserved across streams with naturally varying oxidation activity and methanotroph communities. In addition, we examined intrastream variation in the methanotroph communities, by extracting DNA from six separate gravel samples collected at 50 m intervals along a 250-m reach of the River Lambourn. Details of the DNA extraction method and amplification of *pmoA* gene fragments are provided as Supplementary Information.

Molecular analysis of the *pmoA* gene to assess the natural variation in methanotroph communities

454 sequencing of the *pmoA* gene was performed at the Research and Testing Laboratory (<http://www.researchandtesting.com/index.php>). Sequences were processed in the QIIME pipeline and its associated modules (Caporaso *et al.*, 2010). Details of sequence processing are given in Supplementary Information.

Accession numbers

The DNA sequences reported in this study were deposited in the EMBL database under the accession numbers LK052686-LK052755.

Results

Recovery of ¹³C-CH₄ in ¹³C-organic and ¹³C-inorganic carbon fractions

Methane was oxidised rapidly over the first 20–30 h in each of the eight-batch incubations (Figure 2a) and, although there was some variation in the rates of oxidation, there was no significant change over the entire 409 h of incubation, with an average rate of 1.8 nmol CH₄ g⁻¹ gravel h⁻¹ (±0.2, s.e. *n* = 8). Note,

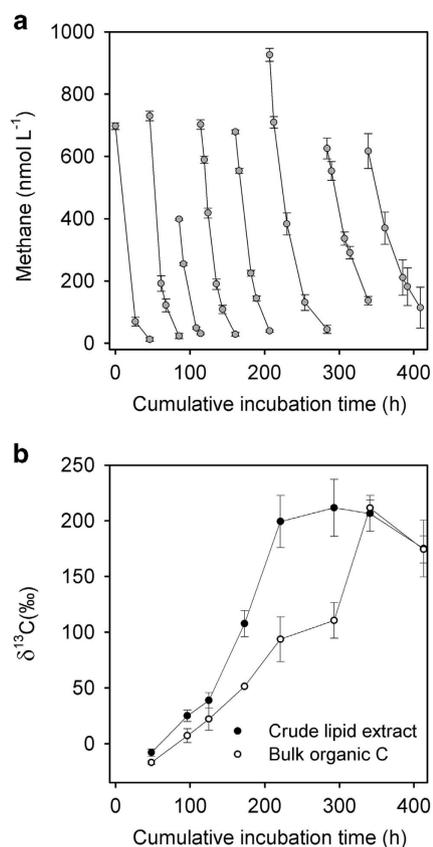


Figure 2 Oxidation and incorporation of ¹³C-CH₄ during repeat batch incubations with riverbed gravel. (a) Rapid initial rates of methane oxidation following each sequential enrichment of the headspace with ¹³C-CH₄. (b) Resulting enrichment of the lipid (crude extract) and bulk organic components of the gravels over the entire 409 h incubation. Mean ± s.e., *n* = 4 for all.

although high at 99 atom% ^{13}C , the degree of ^{13}C labelling used in the experiments had no measurable effect on the rates of methane oxidation (see Supplementary Figure 3). We subdivided the ^{13}C assimilated into a crude lipid extract and bulk organic carbon fraction. The crude lipid extract approached an isotopic steady state (that is, an approximately constant $\delta^{13}\text{C}$) after about 9 days (206 h; Figure 2b), whereas the bulk organic carbon fraction took a little longer. The difference in time to reach isotopic steady state reflected the respective size of each pool, with lipids making up 17% (± 1.2 s.e., $n=8$) of the organic carbon and the bulk fraction 83% (± 1.2 s.e., $n=8$). Although a steady state was only achieved around midway through the eight-batch incubation, there was no significant difference in the split between ^{13}C recovered as lipid or as the bulk organic carbon fraction between each batch (analysis of variance, $F_{6,7}=0.96$, $P>0.05$). Overall, we were able to recover all of the $^{13}\text{C}\text{-CH}_4$ introduced to the vials ($105\% \pm 6$ s.e., $n=8$; one-sample t -test, $t=0.83$, $P>0.05$, see Supplementary Table 1), with it being split (analysis of variance, $F_{7,7}=1.49$, $P>0.05$), on average, between 53% (± 3) inorganic ($^{13}\text{C}\text{-DIC}$) and 52% (± 6) organic carbon (^{13}C -bulk organic plus lipid fractions). Approximately, therefore, for the repeat batch incubations, 50% of methane metabolised was fixed into organic biomass.

Quantifying the CCE by the yield of $^{13}\text{C}\text{-DIC}$

In a follow-up, short (<60 h) incubation, the evolution of $^{13}\text{C}\text{-DIC}$ was linear over the first 28 h, after which the low concentration of $^{13}\text{C}\text{-CH}_4$ limited production of $^{13}\text{C}\text{-DIC}$ (Figure 3a). We then used the ratio of $^{13}\text{C}\text{-DIC}$ produced per $^{13}\text{C}\text{-CH}_4$ oxidised (0.48 ± 0.02 s.e., $n=20$; Figure 3b) during the first 15 h (when linearity was greatest) to estimate CCE through methanotrophy to be 0.52 (that is, $1 - 0.48 - 100 = 52\%$ efficient $\pm 2\%$). Owing to the accumulation of CO_2 from biofilm respiration during these batch incubations (typically $1.6 \mu\text{M h}^{-1}$ in the water phase), the pH declines in parallel at around 0.1 units h^{-1} , which is similar to the diel pattern recorded in these rivers (Shelley *et al.*, 2014); yet, our measure of CCE remains invariant of pH over ~ 20 h. Note that this short incubation estimates of 52% was indistinguishable from that measured directly as fixed ^{13}C -organic carbon over the previous 409 h of repeat batch incubations (52 ± 6 s.e., $n=8$, as above).

The effect of methane concentration on the oxidation kinetics and conversion efficiency

We measured a clear kinetic effect of methane concentration. The rates of oxidation increased some 10-fold over the local seasonal range in river water methane concentration (for example, $\sim 20\text{--}150$ nM, inset Figures 4a and b) and only began to saturate at ~ 7000 nM CH_4 (Figure 4a, main), indicating a high

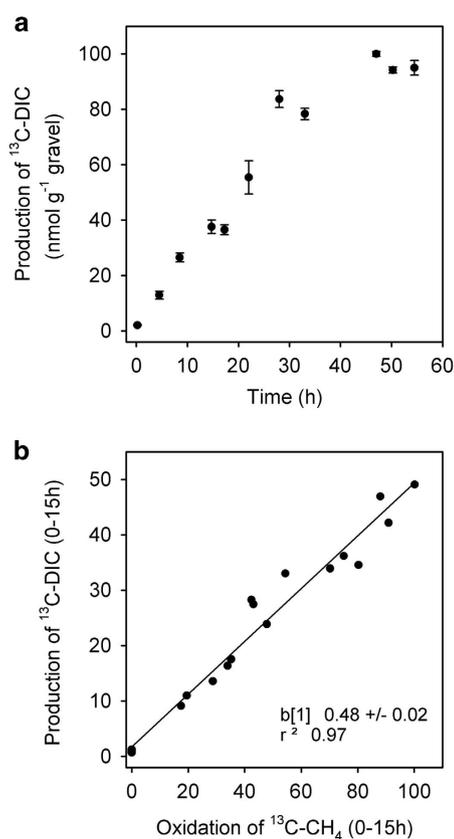


Figure 3 Using the production of $^{13}\text{C}\text{-DIC}$ as a measure of CCE by riverbed methanotrophs during short-term incubations. (a) Production of $^{13}\text{C}\text{-DIC}$ from the oxidation of $^{13}\text{C}\text{-CH}_4$ over ~ 60 h to the point of CH_4 limitation. Mean \pm s.e., $n=5$ (b) as (a) but with the $^{13}\text{C}\text{-DIC}$ as a function of the amount of $^{13}\text{C}\text{-CH}_4$ oxidised during the first 15 h, that is, the period when the linearity of DIC production was strongest ($r^2=0.93$, $P<0.001$) and each datum is the result of a single incubation. Here as the intercept is not different to 0, the slope (0.48) is equivalent to the ratio of DIC to CH_4 and $1-0.48$ is a measure of carbon fixation efficiency, that is, $0.52 \times 100 = 52\%$ fixed. Note this estimate of 52% was indistinguishable from the estimate based on the recovery of organic carbon during the previous repeat batch incubations of 52%.

capacity for methane oxidation in the gravels. In contrast to the marked kinetic effect, there was no statistically significant relationship between the fraction of $^{13}\text{C}\text{-CH}_4$ recovered as DIC and the initial concentration of methane (Figure 4c). Over this range of methane concentrations, the average CCE was 53% (± 1 s.e., $n=40$) and in good agreement with the previous experiments (52%).

Streambed gravels as a natural source of copper

The highest concentrations of copper were measured in the gravel plus UHP water ($7.5 \mu\text{g Cu l}^{-1}$); probably as a consequence of the dissolution of chalk in the mildly acidic UHP water ($\sim \text{pH } 5.5$). Gravels plus stream water yielded a final concentration of $5.1 \mu\text{g Cu l}^{-1}$, which was greater than that for river water only ($4.0 \mu\text{g Cu l}^{-1}$). Although long-term Cu data (see Supplementary Information) are only available for

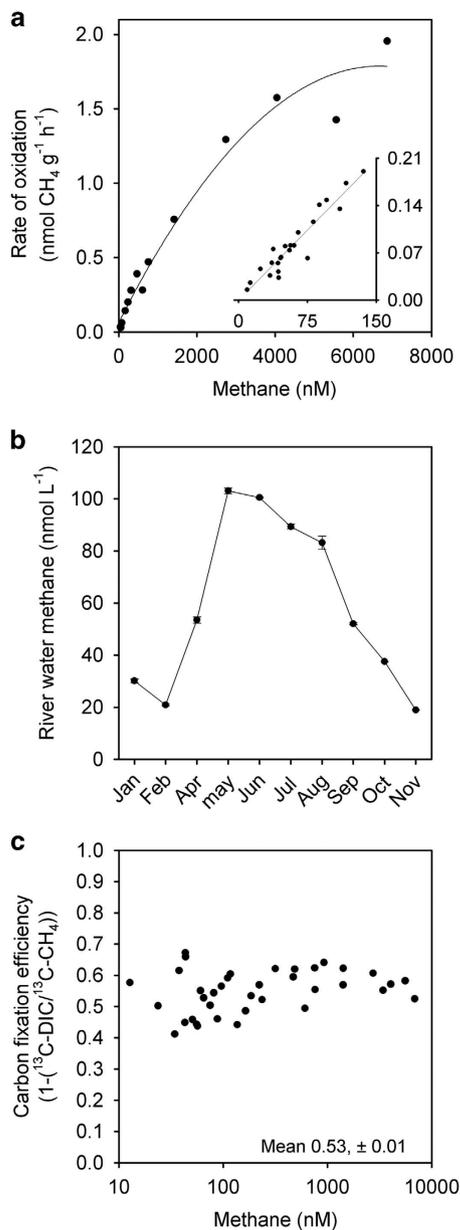


Figure 4 Kinetic effects of methane on its rate of oxidation and the efficiency of carbon fixation. (a) Rate of methane oxidation as a function of methane concentration from above and below ambient river concentrations (insert, first-order linear regression $r^2=0.90$) to the point of near saturation (main panel, second-order polynomial, $r^2=0.97$). (b) Annual pattern of river water methane in the Lambourn (see Shelley *et al.* (2014) for details and Sanders *et al.* (2007) for similar patterns). (c) CCE exhibiting no relationship with methane concentration from ~ 10 nM CH₄ to 7000 nM CH₄. Here the mean CCE value of 53% (± 1) was indistinguishable to that determined either in the repeat batch incubations (52%, ± 6) or that from ¹³C-DIC over 15 h (52%, ± 2 , see Figure 3). Each datum point is the result of a single incubation.

two of the streams visited in the wider survey, the Chess and the Gade, their overall range of 1–5.3 $\mu\text{g Cu l}^{-1}$ agrees with our own measurements in the Lambourn. We would conclude, therefore, that the gravels, and underlying chalk, are a natural source of copper to the chalk streams.

CCE across a natural gradient of variation in methanotroph communities

The gene for particulate methane monooxygenase (*pmoA*) from eight riverbed samples was successfully amplified and then sequenced. A total of 23 212 sequences were clustered into 70 operational taxonomic units (OTUs). Representative sequences of all the OTUs and their closest relatives were aligned and analysed phylogenetically to reveal a diverse methanotroph community, drawn from the eight independent streams (Figure 5a; Supplementary Table 3). A principal coordinates analysis of all the *pmoA* sequences was performed using the Unifrac weighted metric (Figure 5b). This plot indicated that the majority of variation in the methanotroph community (82%) was represented by the first principal coordinate (PC1 in Figure 5b) which was driven by a shift from a predominately Type I community (>98%) at one end of the axis (sample 6, Mimram, Figure 1), to a Type II-dominated community (>90%) at the other (samples 3 and 7, Chess and Cray, Figure 1). Diversity within the eight samples was assessed using the Simpson Diversity Index (Supplementary Table 3), which shows increased diversity in samples between the two extremes, from 0.3 in the Type I-dominated Mimram sample and Type II-dominated Chess and Cray samples, to 0.8 in the more even communities in the Lambourn, Gade and Ver samples (Supplementary Table 3).

Overall though, the distribution of Type I and Type II sequences was almost equal, with a median fraction of total sequence abundance of 42% and 51%, respectively. The majority of the Type II sequences belonged to the *Methylocytis*-related OTU, CS1 (median 48% of all sequences). Although the Type I were best represented (present in all replicates) by the OTU CS2 (median 10% of all sequences), most closely related to clone OK-55 (AB84086). In the Lambourn, we could also assess intrastream variation in community structure using the six samples collected from along the 250 m reach. Here we found greater similarity between the methanotrophs, with a narrower range in PC1 (–0.45 to –0.01, Figure 5b), compared with the samples drawn from across the eight streams (–0.50 to 0.44, PC1, Figure 5b). Although the variation captured by PC2 was not as pronounced as that for PC1, it was similarly greater across the streams, compared with that within the Lambourn.

As expected, gravels from the eight streams all oxidised methane with varying degrees of activity (analysis of covariance, $P < 0.001$, $n = 8$) and there was also marginal variation (analysis of covariance, $P < 0.047$, $n = 8$) in the fraction of ¹³C-CH₄ recovered as ¹³C-DIC (see Supplementary Table 2). In order to derive an overall estimate for the ratio of ¹³C-DIC produced per ¹³C-CH₄ oxidised, while taking into account the moderate difference between the streams, we treated ‘stream’ as a random effect in a mixed-effects model (Zuur *et al.*, 2009). We found

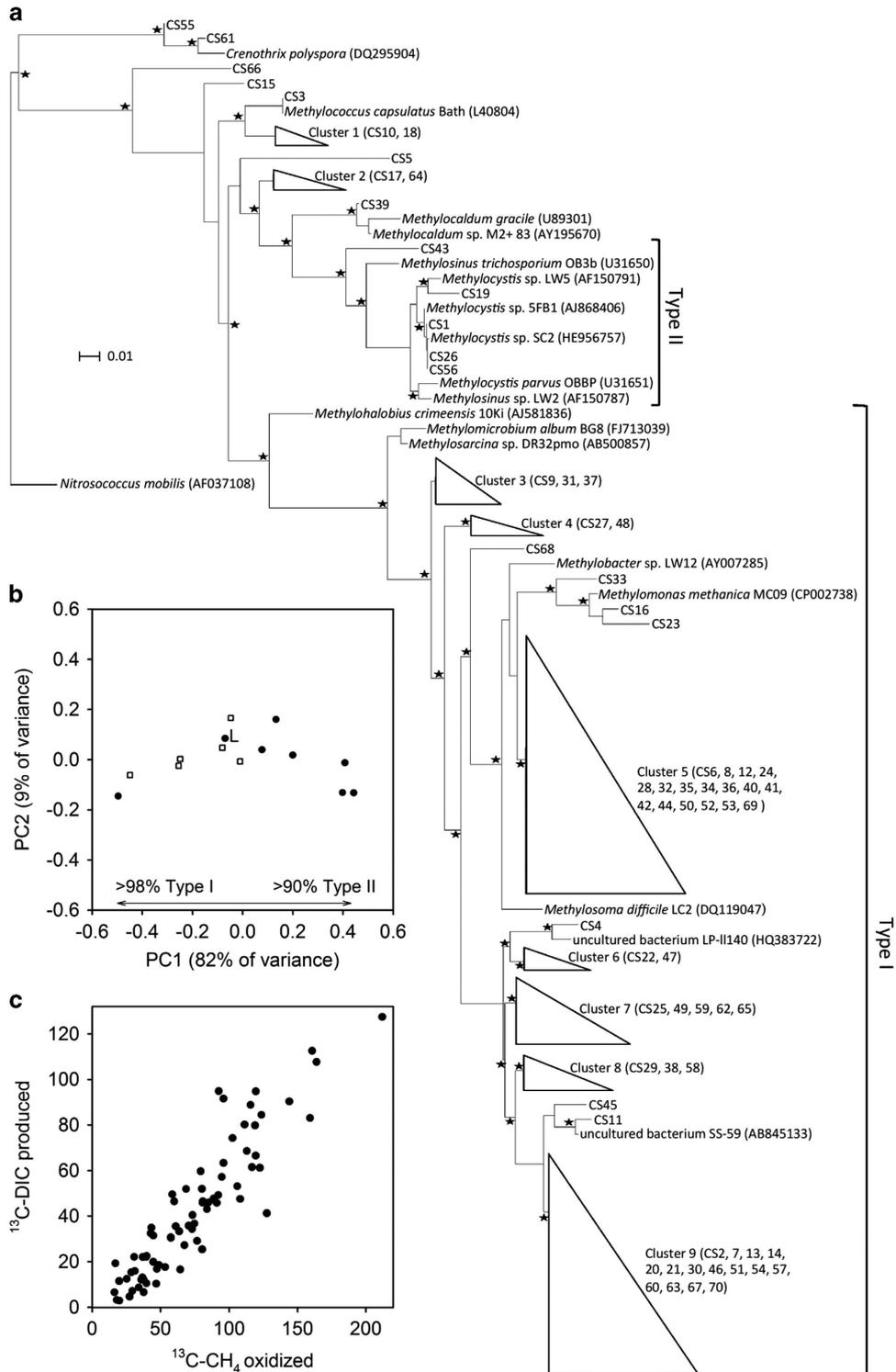


Figure 5 A measure of the variation in the methanotroph community and overall CCE for a sample of eight chalk streams in southern England. **(a)** Maximum likelihood tree of representative *pmoA* sequences (about 200 bp). Collapsed branches are indicated by triangles. The *amoA* gene of *Nitrosococcus mobilis* (AF037108) was used as the out group. Asterisks indicate local support values over 75%. The bar represents 0.01 average nucleotide substitutions per base. **(b)** Principal coordinates analysis of the methanotroph community based on *pmoA* sequences. Filled circles are the eight independent stream samples and open squares the intrastream variation in the Lambourn. L is the pooled sample from the Lambourn also used to measure its CCE. **(c)** Standardized output from the mixed-effect model showing the best estimate of the overall ratio of DIC produced per CH_4 consumed (slope 0.50, s.e. 0.03, t 16.4) and CCE via methanotrophy ($1 - 0.50 \times 100$) was 50% on average for the eight rivers (filled circles in **b**).

that a more complex random intercept and slope model (m1) provided no better a fit to the data than a simpler random intercept only model (m2; analysis of variance m1, m2, $P=0.1203$). The more parsimonious m2, with a fixed slope, fits the data very well (slope 0.50, s.e. 0.03, t 16.4) and, accordingly, CCE ($1 - 0.50 \times 100$) via methanotrophy was estimated to be 50% (95% confidence interval of 44% to 56%) for our sample of eight chalk streams, despite the pronounced variation in the methanotroph community (Figure 5c).

Discussion

Recently, we demonstrated that methanotrophy is widespread throughout the 32 major chalk streams of southern England (Shelley *et al.*, 2014) and that it has the potential to match any increase in methane from methanogenesis that it can intercept before methane escapes the riverbed (Shelley *et al.*, 2015). Here we have clearly shown that the fraction of carbon fixed via that methanotrophy is consistently high which, coupled to the wide range of methane concentrations oxidised by the methanotrophs, will enable methanotrophy to track the seasonal range of methane in the river water as well as exploit the much higher concentrations found in the patches of deposited fine sediments (Sanders *et al.*, 2007; Shelley *et al.*, 2015).

We have used the fraction of ^{13}C recovered as ^{13}C -DIC to more simply quantify the CCE of methanotrophy, without quantitatively extracting and purifying the organic fractions (Maxfield *et al.*, 2012). One could argue that some intermediate metabolite of the methane oxidation pathway (for example, ^{13}C -methanol) could be released during the initial 'linear' phase of ^{13}C -DIC production. This would then be metabolised by non-methanotrophs and, at the same time, the 'loss' could be counterbalanced by fixation of ^{13}C -DIC by other chemototrophs, or via the possible carboxylation reactions of some heterotrophs (Alonso-Saez *et al.*, 2010). If true, we would overestimate the CCE of methanotrophy, but we are confident that such an effect was negligible for the following reasons.

First, the ^{13}C -DIC produced in our experiments was strongly diluted by the ^{12}C -DIC produced by total community respiration, only constituting a maximum of 1.8% of the total DIC pool ($^{12}\text{C}+^{13}\text{C}$ -DIC). Therefore, any contribution of ^{13}C -DIC being fixed by any non-methanotrophic metabolism would have been negligible. Second, there was no increase in the rate of methane oxidation over the 409 h of repeat batch incubations, which indicated no net growth in the population of methanotrophs during this time. In a community at steady state, the potential loss of any ^{13}C -methane would have been small—especially during the short ^{13}C -DIC experiments (Figure 3a). If, however, there was cell death alone, then, not only would the batch rates of methanotrophy have

declined with time, but they would not have reached an isotopic steady state, which was clearly not the case (Figure 2b). Eventually, though, some of the ^{13}C assimilated by the methanotrophs will be reworked and shared among other members of the gravel community, but this process has been shown to be detectable only after some 2–3 weeks in soils (Maxfield *et al.*, 2012) which, again, will not affect our assay over 15 h or less.

We directly determined the CCE to be 50% on average for our sample of eight independent chalk streams. Carbon conversion efficiencies are known to be fundamentally different between Type I and Type II methanotrophs, as the RuMP (and its variants (Anthony, 1982; Trotsenko *et al.*, 2008; Kalyuzhnaya *et al.*, 2013)) pathway of C1 assimilation, largely characteristic of Type I methanotrophs, converts more C into biomass, compared with the serine, Type II, equivalent (Note that Type X methanotrophs contain genes for both assimilatory pathways (Trotsenko *et al.*, 2008)). Pure culture work has also shown CCE and growth rate to be greater when pMMO is expressed rather than sMMO, with that expression of pMMO being strongly regulated by Cu (Leak, Dalton, 1985; Joergensen and Degn, 1987; Semrau *et al.*, 2010). Theoretically, the higher metabolic cost of growing on nitrate could also decrease the CCE compared with growth on NH_4^+ (Leak and Dalton, 1986). Clearly, in a natural mixed community, all of these factors could interact to influence the CCE but, at 50%, our CCE was consistently towards the upper end of the commonly reported range (~30–60% CCE (Leak and Dalton, 1986; Rostkowski *et al.*, 2012; Fei *et al.*, 2014)).

Nitrate provides an abundant (~6 mg l⁻¹ on average) supply of N in both the overlying water and gravelbed of these chalk streams, as it does in many rivers throughout the developed world, (see Supplementary Figure 1 and Pretty *et al.*, 2006; Sanders *et al.*, 2007) whereas, although the riverbed is an area of intense mineralisation, the concentration of NH_4^+ is orders of magnitude lower (Pretty *et al.*, 2006). In addition, the temperature of chalk streams is comparatively stable due to groundwater buffering as per other permeable geologies (for example, sandstone and limestone (Wheater *et al.*, 2007)). Together, the abundant supply of nitrate, along with the chalk acting as a natural source of copper and providing a stable temperature regime, appears to provide a favourable habitat for a variable community of methanotrophs (70 OTUs). Our community represents a substantial proportion of the known diversity for methanotrophs in other aquatic ecosystems (Chen *et al.*, 2008; He *et al.*, 2012; Reim *et al.*, 2012; Yun *et al.*, 2012) and reasserts the view that wherever there is methane and oxygen there will also be methanotrophs (Trotsenko *et al.*, 2008; Semrau *et al.*, 2010).

Most freshwater systems studied to date (Chen and Murrell, 2010) have far higher methane concentrations than our streams (μM – mM relative to nM) and tend to

be dominated by Type I methanotrophs. In contrast, in upland and forest soils, which are typically at equilibrium with the atmosphere (1.8 p.p.m. –80 nmol CH₄ l⁻¹ air), *pmoA* sequences are dominated by the Type II *Methylocystis*. In chalk streams, the prevailing concentration of methane in the gravels is typically 40 to 100 nmol CH₄ l⁻¹ pore-water; although in summer it can reach 200 to 300 nmol CH₄ l⁻¹ in the overlying river water and even 2 to 4 μmol CH₄ l⁻¹ in the anoxic patches of deposited fine sediment (Sanders *et al.*, 2007; Shelley *et al.*, 2015). As expected for the gravelled concentrations of methane, the Type II methanotroph OTU, CS1 (*Methylocystis* related (Figure 5a and Supplementary Table 3)) was a major component of the overall community (median 48% of all sequences). In contrast, however, to the prevailing perspective for freshwater systems and low-methane soils (Chen and Murrell, 2010), we found a more equitable distribution of Type I (median 42% of all sequences) and Type II (median 51% of all sequences) methanotrophs, and there even appeared to be a wider variety of Type I compared with Type II methanotrophs in these low-methane streams. Some have argued that the occurrence of *Methylocystis* sp. SC2, beyond its common habitat of low-methane soils, may in part be due to it encoding, and expressing, multiple methane monooxygenases with substantially different affinities for methane (Baani and Liesack, 2008; Chen and Murrell, 2010). Our observations suggest that such a trait may also be common to a wider group of methanotrophs.

Despite the variation in the methanotroph communities across the stream samples, a consistently high CCE of 50% held for the methanotroph community as a whole. Our findings may suggest a degree of functional redundancy within the variable methanotroph communities inhabiting these streams, although these data do not show which methanotrophs are active in the streams, which would be required to test whether the methanotroph communities in these streams are actually functionally redundant. In addition, some of the detected variation may reflect a suite of *pmoA*-encoded enzymes covering a wide range of affinities for methane (Baani and Liesack, 2008). The latter would certainly make sense given the large range of methane concentrations oxidised here (Figure 4), which would be hard to reconcile with a single enzyme of fixed affinity. Such variation would enable methanotrophs to exploit different niches along the full spectrum of methane concentrations found within rivers, from low concentrations in the clean-gravels, to high concentrations in the deposited patches of fine sediment (Reim *et al.*, 2012).

The annual cycle of temperature in rivers on permeable geologies such as chalk, sandstone and limestone is buffered by the groundwater and might only vary between 8 °C in the winter to 16 °C in summer (Pretty *et al.*, 2006). In addition, we have clearly shown that, at the methane concentrations

typical for the gravels, temperature has no effect on methane oxidation, which is similar to data on temperature effects on methane oxidation in lakes (Nguyen Thanh *et al.*, 2010; Shelley *et al.*, 2015). In contrast, the concentration of methane in rivers can change by several orders of magnitude over the year, reflecting both greater methanogenesis within the river and its wider catchment, plus changes in lateral import and oxidation in the river itself (De Angelis and Lilley, 1987; De Angelis and Scranton, 1993; Kone *et al.*, 2010; Bouillon *et al.*, 2012). Despite the strong kinetic effect of methane concentration on oxidation rate found here, there was no significant effect on the fraction assimilated (that is, the CCE). In contrast to the consistently high CCE we found for streambed gravel communities, the range reported for methanotrophs in the oxic water column of some lakes is extreme (6–77%; Bastviken *et al.*, 2003). In lakes, there appears to be a kinetic effect, with the highest efficiencies associated with low rates of oxidation at low concentrations of CH₄ and vice versa when methane is abundant (Bastviken *et al.*, 2003). Such poor production and limited net growth, at high concentrations, may explain some of the increase in the efflux ratio of methane to carbon dioxide recorded in lakes and wetlands (Yvon-Durocher *et al.*, 2014).

The role of inland freshwaters in the global methane cycle is attracting renewed interest (Bastviken *et al.*, 2011) but whereas the potential for methane oxidation is both high and well characterised in wetlands and lakes (King *et al.*, 1990; Tranvik *et al.*, 2009), data from streams and rivers have been lacking. Here we have demonstrated that the variable riverbed methanotroph communities have both a high capacity to oxidise methane, as well as produce organic carbon with high efficiency. Not only will this help sustain the riverbeds potential to attenuate some of the efflux of methane but, based on the data of Shelley *et al.* (2014), and with a CCE of 50%, we estimate methanotrophy to produce between 2 to 4 g C m⁻² y⁻¹ down to 35 cm into the riverbed. Production by methanotrophy is equivalent to 11% of net benthic photosynthetic primary production at the surface and provides a significant chemosynthetic source of carbon to the stream ecosystem as a whole.

Conflict of Interest

The authors declare no conflict of interest.

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