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Further development and validation of CO₂FOAM for the atmospheric dispersion of accidental releases from carbon dioxide pipelines

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This paper reports on the further development and validation of CO₂FOAM, a dedicated computational fluid dynamics solver for the atmospheric dispersion of Carbon Dioxide (CO₂) from accidental pipeline releases. The code has been developed within the framework of the open source CFD code OpenFOAM® (OpenCFD, 2014). Its earlier version used the homogeneous equilibrium method for fully compressible two-phase flow. Validation of the code against CO₂ releases through vertical vent pipes and horizontal shock tubes was previously reported by Wen et al. (2013). In the present study, the homogeneous relaxation model has been implemented as it is more suited to account for the presence of solid CO₂ within the releases. For validation, the enhanced CO₂FOAM has been used to predict CO₂ dispersion in a range of full scale tests within the dense phase CO₂ PipeLine TRANSPORTATION (COOLTRANS) research programme (Cooper, 2012) funded by National Grid. The test case used in the present study involved a puncture in a buried pipe. The experimental measurements were supplied to the authors after the predictions were completed and submitted to National Grid. Hence, the validation reported here is indeed ‘blind’. The validated model has also been used to study the effect of a commercial building located downstream from the release location.

Key words: Carbon dioxide; Releases from pipeline accidents; Homogeneous Relaxation Model; Blind validation.



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Nomenclature

ρ	Density
U	Velocity vector
P	Pressure
k	Turbulent kinetic energy
ω	Specific dissipation rate
h_s	Sensible enthalpy
I	Identity tensor
σ	Surface stress tensor
Sc_t	Turbulent Schmidt number
Pr_t	Turbulent Prandtl number
g	Acceleration due to gravity
μ	Dynamic viscosity
μ_t	Turbulent viscosity
ν_t	Turbulent kinematic viscosity
β	Mass fraction of gaseous CO ₂
α	Mass fraction of condensed phase CO ₂
T_g	Gas phase temperature
τ	Relaxation time

Introduction

There is growing worldwide interest in Carbon Capture and Storage (CCS). CCS is a technology that prevents large quantities of Carbon Dioxide (CO₂) being released into the atmosphere from the use of fossil fuels in power generation and other industries. The realization of CCS often involves the

54 transportation of compressed CO₂ via high pressure pipelines and process systems often in the dense
55 or super critical phases to ensure efficient high volume transportation capacity (Cooper, 2012). The
56 need has hence arisen to address the potential loss of containment scenarios during the transportation
57 process, as these could pose a risk to people and the environment. Although CO₂ might seem
58 harmless, as it is present in the atmosphere at relatively low concentrations, it can cause serious health
59 risks at higher concentrations (Mahgerefteh et al. 2011). Hence, it is important to assess such risks
60 quantitatively so that appropriate guidelines can be developed to inform the development of pipelines
61 for CO₂ transportation.

62
63 National Grid initiated the COOLTRANS (Cooper, 2012) research programme in 2010 to address
64 knowledge gaps relating to the safe design and operation of onshore pipelines for transporting dense
65 phase CO₂ from large industrial emitters in the UK to storage sites offshore. The University of
66 Warwick was contracted to develop a mathematical model to predict the dispersion in the atmosphere
67 that takes place if a buried pipeline transporting CO₂ is ruptured or punctured. This work forms part
68 of COOLTRANS Work Package 1.4 on “Far Field Dispersion Studies”. The work mainly involved
69 numerical predictions of CO₂ dispersion in the far field taking into account obstacles, ground and
70 gravity effects. It is recognised that for the simulation of the far field dispersion of the released CO₂
71 into the ambient environment, the gravity effect cannot be neglected due to the apparent density
72 difference between the released CO₂ and the ambient air. Turbulence effect is of significant
73 importance for jet dispersion downstream of the rapid expansion zone. The CO₂ jet entrains ambient
74 air and any liquid/solid phase CO₂ at the exit will gradually vaporize/sublimate into gaseous CO₂ due
75 to the mixing with ambient air. As such, the predictions of the correct fluid phase during the discharge
76 and dispersion processes have received special consideration given the very different hazard profiles
77 of CO₂ in the gas and solid states. In the collaborative COOLTRANS project [1], the analysis stages
78 in the numerical study involves: (1) The group from University College London (UCL) to predict the
79 outflow and conditions at the point of pipeline puncture; (2) The group from University of Leeds (UL)
80 to predict the near-field conditions; and (3) Our group at University of Warwick (UW) predicts the far
81 field dispersions using the boundary and initial conditions from UL. The present numerical
82 simulations of the far field dispersion were started using the output from the near-field simulations of
83 UL as boundary and initial conditions (Wareing et al., 2014). The dispersion process is affected by
84 flow induced turbulence which leads to a high mixing rate between the released CO₂ and the ambient
85 air. All these effects have been incorporated into the study along with the influence of wind,
86 topography and obstacles to the releases. The detailed results of these parametric studies will be
87 reported in a later paper although some snapshots are included here to illustrate the effect of obstacles.

88
89 The numerical method presented in this paper to simulate the dispersion of CO₂ in the far field
90 follows several recently published works concerned with CO₂ dispersion. Some of them review by
91 Dixon et al. (2012), most of these but not all are also research done within the context of CCS.
92 Overall, the majority of these studies choose to solve a set of ensemble-averaged, density-averaged
93 forms of the transport equations governing mass, momentum, and energy, involving full buoyancy
94 terms (i.e. non-Boussinesq expression). These equations for averaged quantities are then
95 complemented with a two-equation turbulence models (in most cases, standard k-epsilon). The studies
96 concentrating on near-field especially tend to concentrate on handling the phase transition phenomena
97 associated with accidental CO₂ releases. Wareing et al. (2013a, 2013b, 2014a, 2014b) tested both a
98 Homogeneous Equilibrium model (HEM), and a Homogeneous Relaxation Model (HRM) to handle
99 phase changes of CO₂. Closure of their equation set was achieved via the k-epsilon turbulence model,
100 corrected to account for compressibility effects. Particular attention was given to developing a new
101 equation of state for carbon dioxide (a composite equation using both Peng-Robinson and Span and
102 Wagner equations). Importantly they noticed that their approach is valid, provided the CO₂ dense
103 phase particles are sufficiently small. Their model is well validated against their own experimental
104 observations on high pressure releases of multi-phase carbon dioxide “representative” of medium
105 scale releases arising from an accidental pipeline puncture or rupture. In a similar approach Dixon et
106 al. (2012) also tested both HEM and HRM with lagrangian particles for the solid and the inclusion of
107 water condensation to simulate both horizontal and vertical releases (also using k-epsilon model). In
108 both set of studies, comparison with near-field experimental data showed good agreements. Mazzoldi

109 et al. (2008, 2011) conducted quantitative risk assessment of scenarios involving CO₂ leakages for
110 various CCS projects using PANACHE CFD code. They validated their model using Kit Fox
111 experiments. These researchers focused on the far field results and on provision of an improved
112 understanding of elements constituting the risk associated with an accidental release arising from the
113 transport of the captured CO₂ from the industrial sources to a suitable storage site along high-pressure
114 CO₂ pipelines by constructing hypothetical release scenarios. In both studies, the pipelines were
115 assumed to be placed above ground level, and the downwind distance from the leakage source
116 reached by a harmful concentration level (100,000 ppm or larger) of CO₂ was predicted and used to
117 assess the risk to human health. Mazzoldi et al. (2008) simulated different release scenarios for CO₂
118 using both CFD (based on the Reynolds-averaged Navier–Stokes approach) and Gaussian plume
119 models, and evaluated the predictive performance of these two approaches against two field
120 experiments. They concluded (perhaps not surprisingly) that CFD models offered improved risk
121 assessments for hazards associated with the dispersion of CO₂ clouds compared to the simpler
122 Gaussian plume models. Recently Wen et al. (2013) also used the homogeneous equilibrium method
123 to simulate both horizontal shock tube and vertical vent pipe cases. Uniform atmospheric BCs were
124 assumed and roughness conditions were not taken into account. They also used a standard k-epsilon
125 model. Similarly to previous researchers, good results were obtained when compared with
126 experiments commissioned by National Grid with discrepancies within 25%. Xing et al. (2013)
127 followed the same general approach using FLUENT. They again do not provide density
128 modifications. Their validation is done with a reduced-scale field experiment designed to imitate a
129 CO₂ blowout. Interestingly, they tested three different turbulence models: k-epsilon, RNG and SST.
130 Results were in acceptable agreement with the experimental data (also using Hanna’s statistical
131 analysis) but noted that values found with the RNG model were unsatisfactory. Hsieh et al. (2013)
132 simulated two scenarios: a storage tank release in the vicinity of a cubical obstacle and a pipeline
133 rupture in a complex topography involving two axisymmetric hills. In their approach the density
134 variations of the fluid (containing the dense gas) were simplified using the Boussinesq approximation.
135 It’s not clear whether such approach can be justified for near-field dispersion however, as the
136 molecular weight of CO₂ markedly larger than air. Papanikolaou et al. (2011) conducted numerical
137 simulations of CO₂ release to compare with experimental data taken from the Kit Fox CO₂ gas field
138 experiments. Using ANSYS-CFX, it is not clear though which conservation equations they solve
139 however, turbulence was modelled using the standard k-epsilon. Hedlund (2012) studied the real-
140 world incident involving the catastrophic release of CO₂ using the PHAST package to model the CO₂
141 outburst at the Menzengraben potash mine, where several thousand tons of CO₂ were blown out of a
142 mine shaft. His numerical simulation results showed that in the case of the high momentum release,
143 the leaked CO₂ diluted quickly and never reached the ground surface. In contrast, the low-momentum
144 release resulted in a high concentration gas cloud near the ground surface.

145
146 Regarding initial and boundary conditions, Pontiggia et al. (2009, 2010) showed that better treatment
147 of the upstream ABL, including stability classes, generates much better results for their simulations of
148 Prairie Grass and Falcon tests. Scargiali et al. (2011) applies a k-epsilon model in urban area-style
149 setting for heavy gas dispersion, first involving a stationary pre-release flow field simulation followed
150 by a dynamic after-release flow and concentration field simulations. They used a so-called weakly
151 compressible approximation for the buoyancy treatment instead of the simpler Boussinesq
152 approximation employed elsewhere, in view of the strong density gradients in the proximities of the
153 dense plume. The main hypothesis behind this approximation is that density variations are related
154 only to the mean molecular weight and/or temperature changes in the fluid, while density is assumed
155 to be independent of the pressure field, they thus assumed a reference pressure independent of density.
156 Using chlorine, they found that the presence of buildings reduces the maximum ground concentrations
157 while enlarging the affected area. Due to the larger negative buoyancy effects, increasing the amount
158 of heavy-gas released slows down the cloud and increases (normalised) maximum concentrations and
159 lateral spread of the cloud. While it proves quite difficult to establish general trends in the predictive
160 abilities of the aforementioned modelling approaches. On the positive side, they are all capable (for all
161 heavy gases) to predict the trends of the experiments within acceptable level of accuracy. Each
162 provides a detailed picture of the flow and its interaction with the domain. They also serve as useful
163 guides to what modelling approach to adopt.

164 **Mathematical formulation and numerical framework**

165 CO₂FOAM, a dedicated solver for CO₂ dispersion has been developed within the frame of the open
166 source computational fluid dynamics code OpenFOAM® (OpenCFD, 2014). The code allows
167 simulations with the relatively more efficient Steady or Unsteady Reynolds Averaged Navier-Stokes
168 (RANS or URANS) approach as well as the more robust but computationally more intensive large
169 eddy simulation (LES) approaches. It is known that the RANS approach which solves the time/density
170 averaged transport equations is limited by the turbulence models. The LES approach aims at full
171 resolution of the large, energy-bearing structures to momentum and energy transfer and only uses sub-
172 grid scale models for the small eddies not resolved by the computational grids. Since the small scales
173 tend to be more isotropic and universal in nature, their modelling is expected to be more amenable to
174 success and require fewer adjustments when applied to different flows than models in the RANS
175 approach. However, as the use of the LES approach will be computationally too expensive for the
176 present study which involves relatively large computational domains and long durations, the majority
177 of the simulations reported are conducted with the RANS approach with the k- ω SST turbulence
178 model. In order to assess the potential loss of accuracy, preliminary predictions have been carried out
179 with both approaches and found to show similar level of agreement with the experimental
180 measurements. However, the LES predictions have captured the continuing rise in concentration after
181 the cloud arrives at the probe whereas the RANS approach predicted an increasing gas turbulent
182 diffusion. This suggests that the effect of density interface instability and turbulence damping are
183 better captured by the LES. Quantitatively, the comparison shows that the differences in the predicted
184 CO₂ concentration levels are only noticeable within the first 80 m from the release point, further from
185 there, the RANS predictions are in reasonably good agreement with the mean values predicted by the
186 LES approach. This has established the confidence in using the RANS approach when only the far
187 field results of importance. The computational domain extends up to 250 m in the direction of the
188 wind, capturing the far field dispersion.

189
190 All dense gas dispersion and heat transfer phenomena need to be included in atmospheric dispersion
191 models used for predicting the dispersion of a CO₂ cloud formed from a pressurised release.
192 Specifically, these include variable temperature and density, an appropriate turbulence model for
193 dense gas dispersion, an adequate ground-level heat transfer model and gravity flow equations. It is
194 also necessary to include atmospheric stability, complex terrain and source conditions. In order to
195 reproduce field experiments for validation it may also be necessary to include time dependent wind
196 speed and direction. The specification of appropriate inflow and surface boundary conditions, material
197 properties and an appropriate turbulence model are hence crucial to the reliability of the predictions.
198 The turbulence models, in particular, should include components which adjust for the damping effects
199 of thermal or density stratification on turbulent generation and dissipation.

200
201 In the light of previous studies and experimental evidence (Mazzoldi et al. 2011 and Wareing et al.
202 2013), the approach undertaken by UW is in the first place to use the near-field predictions of
203 Wareing et al. (2014) as the boundary and initial conditions for the far field simulations. Following on
204 from the analysis conducted by Weber (2011), both the Homogeneous Equilibrium Method (HEM)
205 and the Homogeneous Relaxation Model (HRM) should be applicable to the present applications. The
206 earlier version of the CO₂FOAM code used the HEM approach to handle the fully compressible two-
207 phase flow. Its validation against CO₂ release tests through vertical vent pipes and horizontal shock
208 tubes was reported by Wen et al. (2013). In the HEM approach, the constituent fluid phases are
209 assumed to remain at thermal and mechanical equilibrium during the decompression process.
210 Consequently, phase-slip and non-equilibrium liquid/vapour transition phenomena, such as delayed
211 bubble formation are ignored (Zucrow and Hoffman, 1975).

212
213 The HRM approach has been implemented into CO₂FOAM in the present study. In this approach, the
214 assumption of mechanical equilibrium, i.e. no phase slip, is retained. However, non-equilibrium
215 liquid–vapour transition is accounted for by a relaxation to thermodynamic equilibrium following
216 Downar-Zapolski et al., (1996) and Brown et al., (2013). The choice of the time constant for the HRM
217 also follows that of Brown et al. (2013). Although the pressure is lower than the CO₂ triple point

218 pressure at the immediate vicinity of the release point, the framework assigned to Warwick team is the
 219 atmospheric dispersion away from the three-phase, quickly expanding stages in the near proximity of
 220 the release from the pipeline. The HRM is hence adequate to model the mixture of solid and gas in
 221 such a far field dispersion study. The sizes of the solid CO₂ particles are assumed to be small enough
 222 so that modelling the CO₂ phase change as relaxation towards equilibrium composition holds. The
 223 deposition or rain out of dry ice is not handled in this model in the light of the reported inconsistent
 224 presence of deposited ice in experiments (GL, 2011).

225
 226 Following previous researchers (Wen et al., 2013 and Hsieh et al., 2013), the RANS approach has
 227 been used to solve the three-dimensional conservation equations for a CO₂/air mixture for mean
 228 (ensemble-averaged) quantities in a turbulent flow field and are solved using governing equations
 229 being shown in the following section.

230 231 **The Governing Equations**

232 The compressible flow conservation equations are Favre averaged (mass weighted), any flow quantity
 233 f are split into mean and fluctuating component as $f = \bar{f} + f''$ with $\bar{f}'' = 0$ and $\bar{f} = \frac{\bar{\rho}f}{\bar{\rho}}$. The (\bar{f})
 234 overbar quantity represents Reynolds averaged mean quantity.

235
236 Continuity

$$237 \quad \frac{\partial \bar{\rho}}{\partial t} + \nabla \cdot \bar{\rho} \tilde{U} = 0 \quad (1)$$

238 Momentum

$$239 \quad \frac{\partial \bar{\rho} \tilde{U}}{\partial t} + \nabla \cdot \bar{\rho} \tilde{U} \tilde{U} = -\nabla \bar{P} - \nabla \cdot \sigma + \bar{\rho} g \quad (2)$$

240 Where

$$241 \quad \sigma = (\mu + \mu_t) (\nabla \tilde{U} + \nabla \tilde{U}^T) - \frac{2}{3} (\mu_t \nabla \cdot \tilde{U} I + \bar{\rho} \tilde{k}) \quad (3)$$

242 In the treatment of buoyancy forces, the large density variations expected from the presence of dry ice
 243 and the molecular weight of CO₂ means that the Boussinesq approximation often applied in
 244 atmospheric dispersion studies cannot be applied (Hsieh et al., 2013). To model the
 245 diffusivity/turbulent diffusion flux, by analogy with Fick's law, a gradient law is employed. The
 246 transport equation for the mass fraction of gaseous CO₂ - ' β ' can be written as:

$$247 \quad \frac{\partial \bar{\rho} \tilde{\beta}}{\partial t} + \nabla \cdot \bar{\rho} \tilde{U} \tilde{\beta} = -\nabla \cdot \frac{\mu_t}{Sc_t} \nabla \tilde{\beta} - \tilde{\alpha} \frac{P_v - P_s}{\tau p_s} \quad (4)$$

248 A simple sub-model for the relaxation to equilibrium has been included, in which the temperature
 249 relaxation is ignored and the condensed phase mass fraction is simply given by:

$$250 \quad \frac{\partial \bar{\rho} \tilde{\alpha}}{\partial t} + \nabla \cdot \bar{\rho} \tilde{U} \tilde{\alpha} = -\nabla \cdot \frac{\mu_t}{Sc_t} \nabla \tilde{\alpha} - \tilde{\alpha} \frac{P_v - P_s}{\tau p_s} \quad (5)$$

251 where p_v is the vapour pressure, p_s is the saturation pressure and τ is a relaxation time taking value of
 252 0.1 ms. Taking species mass diffusion into account, the turbulent viscosity (μ_t) is corrected with the
 253 turbulent Schmidt number (Sc_t).

254
 255 Previous research on atmospheric dispersion has used a range of values for Sc_t between 0.3 and 1.4
 256 depending on various factors including presence of obstacles. In the first instance a value of $Sc_t = 0.7$
 257 was selected following Hassid (1983) for neutral atmospheric conditions. A default turbulent Prandtl
 258 number of 0.85 is applied.

259 260 **Energy Equations**

261 Finally, energy conservation in the system is considered through the sum of sensible-enthalpy and
 262 turbulent kinetic energy equation.

$$\frac{\partial \bar{\rho}(\tilde{h}_s + \tilde{k})}{\partial t} + \nabla \cdot \bar{\rho} \tilde{U}(\tilde{h}_s + \tilde{k}) = \frac{D\bar{P}}{Dt} + \nabla \cdot \frac{\mu_t}{Pr_t} \nabla \tilde{h}_s + \nabla \cdot \frac{\mu_t}{Sc_t} \nabla \tilde{\beta}(\tilde{h}_{CO_2} - \tilde{h}_{air}) + Q_{sub} \quad (6)$$

where, h_s is the sensible enthalpy of the fluid $h_s = \sum_i Y_i h_i$, is the energy flux, Q_{sub} represents the volumetric heat sources for the heat transfer due to sublimation of solid CO₂ particles. The following assumptions apply to the study of the dispersing jet: the effect of body forces (i.e. Coriolis forces) has no relevance on the system energy. Energy dissipation due to viscous forces in gas flow under a turbulent regime may be neglected; energy sources due to compressibility effects are ignored, since large pressure differences are not expected inside the cell; and the heat flux accounts for the heat conduction and the heat flux due to species diffusion with different enthalpies.

The JANAF thermochemical tables¹ (NIST, 1990) are used to calculate the specific heat at constant pressure and the sensible enthalpy for each species. The temperature and density are calculated assuming a homogeneous mixture of ideal gases. The Sutherland (1983) model is used to account for the temperature dependence of the viscosity of each specie in the mixture:

$$\mu(T_g) = \frac{A_s \sqrt{T_g}}{1 + T_s / T_g} \quad (7)$$

where the Sutherland-law constants, A_s and T_s , are obtained from the JANAF database (NIST, 1990).

The system of partial differential equations was solved numerically using a collocated, finite-volume method. Diffusive volume-face fluxes were discretized using a second-order accurate central differencing scheme. The advective volume-face fluxes were approximated using a second-order accurate limited linear scheme. The transient term was discretized using a fully implicit, second-order accurate three-time-level method described in Ferziger and Peric (1996). The PIMPLE algorithm was used to combine the momentum and pressure equations (OpenCFD, 2014). PIMPLE (PIso-SIMPLE) algorithm is merger of semi-implicit method for pressure linked equations (SIMPLE) by Patankar (1980) and pressure implicit with splitting of operation (PISO) by Issa (1986), SIMPLE algorithm in outer iteration is merged with the PISO algorithm as inner iterations to correct the velocity and pressure explicitly.

Turbulence modelling

A turbulence model is required to close the set of averaged transport equations in RANS simulations. A wealth of studies examining the different choice of turbulence models for the simulation of gas dispersion in the atmosphere with or without obstacles or complex terrains exist. Several models, namely the k- ω shear stress transport (SST), k- ϵ RNG and realizable k- ϵ have shown better results in different cases with some performing better than others in various conditions. However, these models are often tested for handling high turbulence flow and in particular separating flow around buildings. While Wareing et al. (2014) employed a compressibility-corrected k- ϵ model in their numerical simulations of the near field flows, no comparative studies have been reported for applications that are similar to the present study for all these models. Hence the previous recommendations need to be treated with caution in evaluating the suitability of the models for CO₂ dispersion.

In the present case density gradients are practically entirely related to concentration gradients, as thermal gradients contribution due to Joule-Thomson effects and atmospheric conditions are negligible in the relatively short domain. The model chosen must be able to capture correctly the phenomenon of gravity slumping associated with dense gas dispersion. Standard two-equations turbulence models, initially developed for constant-density flows, do not account for the turbulence generation or suppression due to buoyancy forces. Thus, to account for this variation in the density and production of turbulence due to buoyancy, a source term is added to both equations. The source

¹ The JANAF thermochemical tables are a comprehensive thermodynamic database for pure substances.

311 term is modelled by either the simple gradient diffusion hypothesis (SGDH) or the generalized
 312 gradient diffusion hypothesis (GGDH) (Daly et. al., 1970). The SGDH is based on the standard
 313 Boussinesq gradient diffusion concept. The main difference between the two methods is the inclusion
 314 of the cross-stream density variation in GGDH, whereas the SGDH models only include the
 315 streamwise density gradient for a vertical plume. The SGDH is found to under-estimate buoyancy-
 316 produced turbulence and offer little improvement from the standard SST formations while the GGDH
 317 was found to offer improved predictions (Kumar and Dewan, 2014 and Maele and Merci, 2006). To
 318 further evaluate the potential improvement in the predictions which can be achieved by the use of
 319 GGDH and improved wall functions, e.g. Apsley (2007) extended to arbitrary wall roughness.

320

321 In order to better capture the interaction between the dispersed CO₂ and the atmospheric specific
 322 boundary layer (ABL-specific), a compressible form of the k- ω SST turbulence model (Menter, 1994)
 323 is used in conjunction with ABL-specific wall-functions for turbulence modelling in the RANS
 324 approach. The SST model (Menter, 1993) is essentially a two-equation eddy-viscosity model. It
 325 combines the best of the k- ϵ and baseline k- ω models with ω being the turbulent frequency (Menter,
 326 1993). The use of the k- ω formulation in the inner parts of the boundary layer makes the model
 327 directly usable all the way down to the wall through the viscous sub-layer. Hence, the SST k- ω model
 328 can be used as a low Reynolds number turbulence model without any extra damping functions. The
 329 SST formulation also switches to k- ϵ type behaviour in the free-stream and thereby avoids the
 330 common k- ω problem that the model is too sensitive to the inlet free-stream turbulence properties. A
 331 blending function F_1 is used that depends on the wall distance and flow conditions. The equations for
 332 k and ω are defined as follows:

$$333 \quad \frac{\partial \bar{\rho} \tilde{k}}{\partial t} + \nabla \cdot \bar{\rho} \tilde{U} \tilde{k} = \nabla \cdot (\mu + \sigma_{k1} \mu_t \nabla \tilde{k}) + P_k - \beta' \rho \tilde{k} \tilde{\omega} \quad (8)$$

$$334 \quad \frac{\partial \bar{\rho} \tilde{\omega}}{\partial t} + \nabla \cdot \bar{\rho} \tilde{U} \tilde{\omega} = \nabla \cdot (\mu + \sigma_{\omega 1} \mu_t \nabla \tilde{\omega}) + 2(1 - F_1) \sigma_{\omega 2} \rho \frac{1}{\omega} \nabla \tilde{k} \cdot \nabla \tilde{\omega} + \frac{\alpha_3}{v_t} P_k - \beta_3 \rho \tilde{\omega}^2 \quad (9)$$

335

336 with β' , α_1 , σ_{k1} , $\sigma_{\omega 1}$, $\sigma_{\omega 2}$, α_3 and β_3 , as empirical constants (Menter, 1994). Due to the baseline
 337 formulation that does not account for the transport of turbulent stress, the eddy-viscosity would be
 338 over-predicted when calculating the turbulent viscosity from k and ω . A limiter to the kinetic viscosity
 339 is introduced to obtain such transport behaviour:

$$340 \quad v_t = \frac{a_1 \tilde{k}}{\max(a_1 \tilde{\omega}, SF_2)} \quad (10)$$

$$341 \quad \text{with} \quad \mu_t = \rho v_t \quad (11)$$

342

343 F_2 is a blending function similar to F_1 that restricts the limiter to the wall boundary layer and S is an
 344 invariant measure of the strain rate. More details especially on the blending functions F_1 and F_2 can be
 345 found in (Menter, 1993).

346

347 As observed by Ayrault et al. (1998) in studying the effect of negative buoyancy on plumes, the
 348 phenomenon of gravity slumping is closely associated with dense gas dispersion. They found that
 349 although the post-obstacle dispersion behaviour of the plumes mimicked that of a passive gas, the
 350 spread and effect on gas concentration were clearly marked on the upwind side and in-between the
 351 obstacles. They also found that the turbulence and concentration statistics showed less root mean
 352 square (RMS) variations than for neutral conditions, indicating the damping effect of density gradient
 353 on turbulence and dispersion. In order to capture this effect, it is important to incorporate appropriate
 354 buoyancy modifications. Thus, to account for this variation in the density and production of
 355 turbulence due to buoyancy, a source term is added to both equations. The source term is modelled by
 356 two methods:

357

1. The simple gradient diffusion hypothesis (SGDH)

358
$$S_B = -\frac{\mu_t}{g} \frac{1}{\bar{\rho}^2} \nabla \rho \cdot (\nabla P + \rho_\infty g)$$

359 2. The generalized gradient diffusion hypothesis (GGDH).

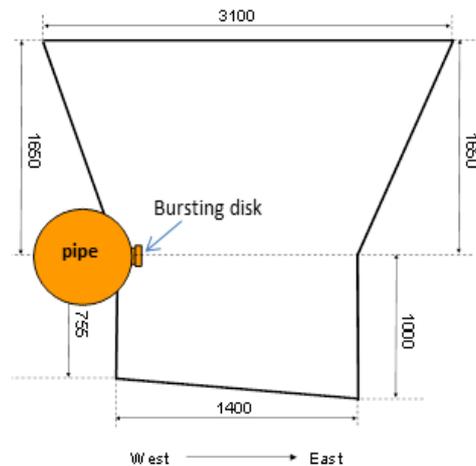
360
$$S_B = -\frac{3}{2} \frac{\mu_t}{g} \frac{1}{\bar{\rho}^2 k} (u'u' \nabla \rho) \cdot (\nabla P + \rho_\infty g)$$

361 The former is found to under-estimate buoyancy-produced turbulence while improved predictions
 362 were reported with the latter (Kumar and Dewan, 2014 and Maele and Merci, 2006). However, little
 363 information was found on the performance of these models in negative buoyancy situations. The
 364 SGDH model was used by previous authors (Scargiali et al., 2011) simulating heavy gas dispersion
 365 with non-negligible effects. It was hence decided to test both methods.

366
 367

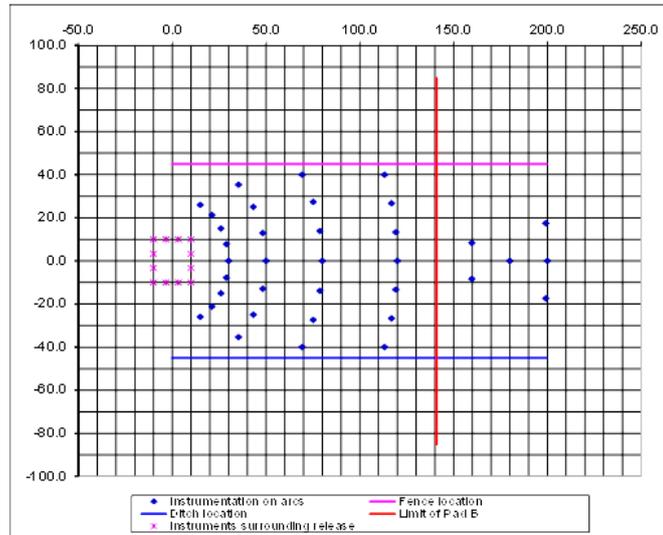
368 **Experiments considered**

369 The experiment considered is Case Study 3 in the COOLTRANS research programme. It involved a
 370 nominally steady release from the side of a length of horizontal, below ground pipe, that was kept
 371 filled with CO₂ in the dense phase. The experiment was carried out to study how a puncture in a
 372 below ground pipeline would behave. The test section had a bursting disk fitted to one side. This disk
 373 failed at a pressure of approximately 150 barg, releasing CO₂ into a pre-formed crater below ground.
 374 The rig was instrumented with pressure transducers and thermocouples and the pressure and flow rate
 375 in the charge line were measured. The configuration resulted in a quasi-steady flow through the
 376 opening on the side of the pipe into the pre-formed crater. Approximately 30 kg/s of CO₂ left the
 377 crater at a maximum velocity of around 23 m/s. Measurements were made of the concentration and
 378 temperature within the resulting CO₂ cloud, as it dispersed in the atmosphere. A cross-section of the
 379 pre-formed crater that was manufactured to reproduce the crater formed in one of the earlier puncture
 380 experiments are shown in figure 1 and the locations of the temperature and CO₂ sensors are shown in
 381 Figure 2. The release conditions are summarised in Table 1.



382
 383 Figure 1. A cross-section (Elevation) of the preformed crater through the point of release constructed
 384 to surround the release location in Case Study 3 of the COOLTRANS research programme,
 385 dimensions in mm. (reproduced form Allason et. al., (2012)).

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Figure 2. Locations of external temperature and concentration measurements in Case Study 3 of the COOLTRANS research programme (reproduced from GL (2011)).

Table 1: Initial Conditions for the puncture test (Reproduced from Wareing et al., (2014))

Section	Item	Value		
Release	Diameter of release	25	mm	
	Outer diameter of test section	914	mm	
	Wall thickness of test section	25.4	mm	
	Depth of top of test section below local ground level	1.2	m	
	Orientation of release	Horizontal at mid height of test section		
	Gas composition	Component	Mole %	
CO ₂		100		
Atmospheric and external conditions – average value in 30 seconds prior to test given	Temperature	3.6	°C	
	Relative humidity	81	%	
	Average of wind speed measured at	0.73 at 4m	m/s	
		0.61 at 1m		
	Average of wind direction measured at	255 at 4m	°	
260 at 1m				
Surrounding terrain local to the point of release	Nominally flat			

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Computational Set-up

In order to carry out full scale simulation of the field experimental trial, the crater inlet, atmospheric boundary conditions (wind) and ground conditions all need to be included in the model. The computational domain orientation is shown in figure 3.

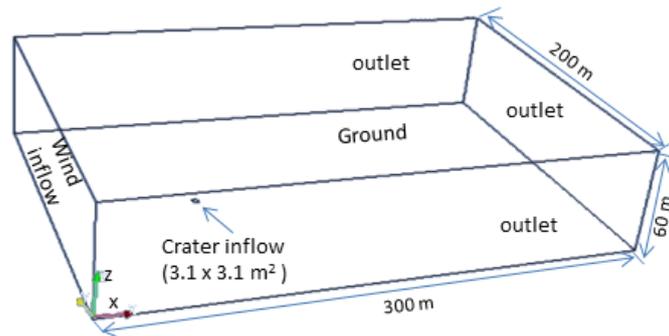


Figure 3. Computational domain

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Crater inlet

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403 In practice, risk analysts dealing with complex systems may have to apply a chain of numerical
404 models, where the output of the first computation is used as input to the next one, etc. In this case, we
405 can identify three particular release processes: in the pipeline, crater and atmosphere. As planned, the
406 far field dispersion study has drawn upon the near-field dispersion studies by researchers at UL to
407 provide inlet conditions while the computations of UL used the decompression simulations of the
408 group from University College London. More specifically; the crater inlet data were taken from a
409 concurrent study by Wareing et al. (2014), who simulated specifically the exit flow inside the crater
410 corresponding to the rupture experiment. The UL's near-field predictions sampled from a plane at $H =$
411 0 was interpolated onto the computational domain in the present study to provide adequate boundary
412 conditions for the CO_2 release from the crater. The crater inlets is placed at the ground level 80 m
413 from upstream wind inlet and correspond to the origin of the frame. An unstructured mesh
414 containing approximately 2.8 M cells with finer resolution around the crater inlet and the obstacle
415 and coarser elsewhere.

416
417 The experiment showed that a high momentum jet of a mixture of air and solid/gas CO_2 escapes the
418 crater. This highlights the importance of resolving this CO_2 source as opposed to using a point or
419 simple area source often used in previous studies (Puttock, 1987). The released cloud carries
420 significant momentum and this jet phase dominates its early dispersion, concurrently the air flow
421 surrounding the crater is expected to be strongly affected by the release creating local recirculation
422 and perturbing the equilibrium parabolic velocity profile set at the wind inlet. The details of this wind
423 inlet are given in the following section. Another feature is the concentration of gas and solid CO_2 at
424 the inlet. It is clear that a great deal of mixing took place within the crater and that the original dense
425 phase CO_2 has expanded, resulting in a mixture of gaseous and solid CO_2 ejecting from the crater
426 (Dixon et al. 2012, Mazzoldi et al. 2008 and Wareing et al. 2014).

427
428 The impact of the inlet conditions constitutes a certain departure from most other toxic gas dispersion
429 studies in that the validity of the computational fluid dynamics (CFD) model is judged both on its
430 ability to deal with the near-field and the more passive dispersion further outwards. Also important in
431 the computational set-up is the specific toxicity of CO_2 . Recent interest in CCS technology has
432 prompted the reconsideration of the potential hazard of CO_2 in the context of potential very large
433 release during CCS scale operation, which has the potential to produce a harmful effect (Engebø et al.
434 2013 and McGillivray et al., 2014). An important feature here is that the crater boundary also "sucks
435 in" air from the atmosphere. When the flow thus exits this domain through the crater, zero gradient
436 boundary conditions are set for the pressure.

Atmospheric and terrain Conditions

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439 An aerodynamic surface roughness length of about 1 cm was determined to give the best fit to a
440 logarithmic velocity profile, assuming a neutral stability atmosphere. That is the value of z_0 in the
441 expression for the wind velocity profile below:

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$$u(z) = \frac{u^* \log\left(\frac{z}{z_0}\right)}{k} \quad (12)$$

The atmospheric boundary layer at the inflow boundary is characterized by a velocity of 0.75 m/s at a 4 m height following a log-law (constant velocity boundary condition). The surface roughness length z_0 is 1.74 cm (chosen according to the terrain vegetation which is grass and bush covered) and the friction velocity u^* is 0.8 m/s. The ambient air temperature is 283.15 K and humidity is 80%.

In order to model the atmospheric boundary layer (ABL), the inflow boundary conditions were chosen following the approach recommended by Richards and Hooxey (1993), Blocken et al. (2007) and Parente et al. (2011). This involved applying consistent boundary layer profiles for velocity corresponding to a neutral ABL (as indicated by the pressure gradient scheme). To be consistent with the inlet, the top boundary is set as a constant shear boundary. This implies that the top boundary should be high enough to be little disturbed by the crater release and/or obstacles.

Computations were conducted firstly to predict the steady-state atmospheric flow field with the area of the crater inlet treated as the ground. This flow field was then used as the initial condition for the subsequent predictions of the dense gas dispersion, which occurred when the gas mixture inside the crater was released to the atmosphere (time $t = 0$ s) by changing the crater boundary to CO_2 inflow. The interval of each time step was automatically adjusted during the computation to satisfy the Courant–Friedrichs–Lewy (CFL) condition of 0.1.

Wind direction is generally non-stationary and depending on meteorological conditions, may vary widely within a short period of time, i.e. a few minutes, in other words of the same order as an accidental release. Most previous CFD studies used a constant mean wind direction at inlet. However, this assumption as observed by Hanna et al. (1993), have led to previous RANS simulations of atmospheric dispersion showing a large discrepancy in lateral spread between simulations and experimental data. It also caused a significant overestimation of the concentrations in a vertical plane through the point of release and parallel with the wind direction. On the other hand, accounting for the full wind direction variability resulted in over-prediction of the lateral spreading of pollutants in some previous studies. It is, hence, thought that the measurements, which themselves are subject to the sensor response time and the actual readings, lie in between the constant wind and the variable-wind solutions.

Results and validation

Grid sensitivity study

The dimensions of the computational domain are 300 m (L) by 200 m (W) by 60 m (H). The X-axis is in the horizontal streamwise direction, i.e. parallel to the wind direction. The Y-axis is lateral and perpendicular to the wind direction while the Z-axis is in the vertical direction (shown in Figure 3). The CO_2 /air mixture is released from the crater inlet at an approximate mass flow rate of 50 kg/s. All other sides (both lateral and downwind side), zero-gradient boundary conditions were chosen for the CO_2 mass fraction.

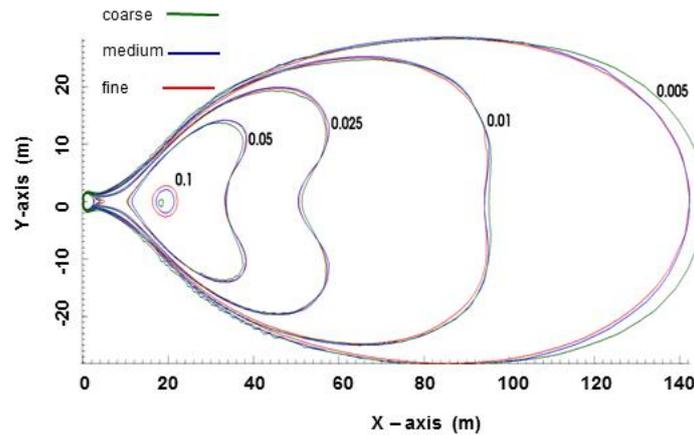
A grid sensitivity study has been conducted with three different mesh resolutions. The coarse, medium and fine meshes include 880,500, 2,577,040 and 5,792,900 grid points; respectively as shown in Table 2. Figures 4 and 5, illustrate that the predictions by the medium and fine meshes are almost identical. For the medium mesh the typical computational time was around 40 hours using 72 processor cores.

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Table 2 Mesh attributes tested

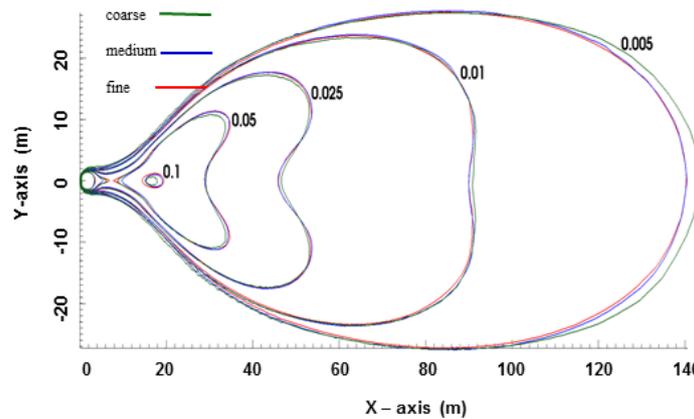
	Points on inlet circle	Minimum dimensions (m)			Maximum dimensions (m)			Number of cells
		x	y	Z	x	y	z	
Coarse	50	0.20	0.20	0.30	4.58	4.00	2.35	880,550
Middle	80	0.12	0.12	0.25	3.23	2.85	1.85	2,577,040
Fine	100	0.098	0.098	0.20	3.05	2.22	1.34	5,792,900

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Figure 4. Comparison of the predicted CO₂ contours at ground level using different grid resolutions.



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Figure 5. Comparison of the predicted CO₂ contours at 1m above ground using different grid resolutions.

505 In Table 3, the predictions with different grid resolutions for nine monitoring points are further
 506 compared, location of monitoring points are shown in Figure 6. It is seen that the largest differences
 507 between the predictions of the fine and medium resolutions is 6% for one point while for all other
 508 monitoring points the differences between the two sets of predictions are generally less than 2%. It
 509 was hence decided that the medium grid resolution will be used for the subsequent validation and
 510 parametric studies to investigate the effects of obstacles, slopes, wind speeds and directions. The first
 511 spacing is about 0.04 m and the grid size on the surface varies from 0.5 m close to the crater to 30 m
 512 in the far field. The maximum cell aspect ratio is 30, the maximum grid skewness is 1.2 and the mesh
 513 non-orthogonality is 10 (averaged) with a maximum of 40.

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Table 3 Summary of the grid sensitivity study in percentage differences

Points	β at 3 m height			$\frac{\beta_{coarse} - \beta_{middle}}{\beta_{middle}} \%$	$\frac{\beta_{fine} - \beta_{middle}}{\beta_{middle}} \%$
	coarse	middle	fine		
A	0.062	0.063	0.062	-0.092	-2.06
B	0.014	0.013	0.014	+6.92	+6.02
C	0.016	0.016	0.016	+2.24	+0.18
D	0.018	0.019	0.018	-3.45	-1.67
E	0.010	0.010	0.010	+1.60	+0.43
F	0.007	0.007	0.007	-2.14	-1.09
G	0.007	0.007	0.007	+1.00	+0.44
H	0.007	0.007	0.007	+0.63	-0.01
I	0.005	0.005	0.005	+0.47	+0.45

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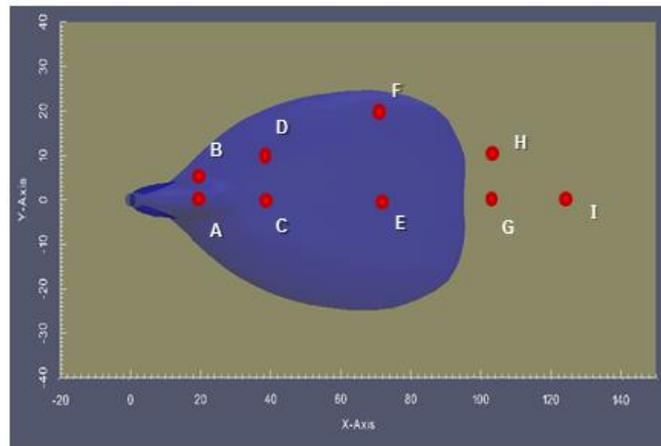


Figure 6. Monitoring point locations

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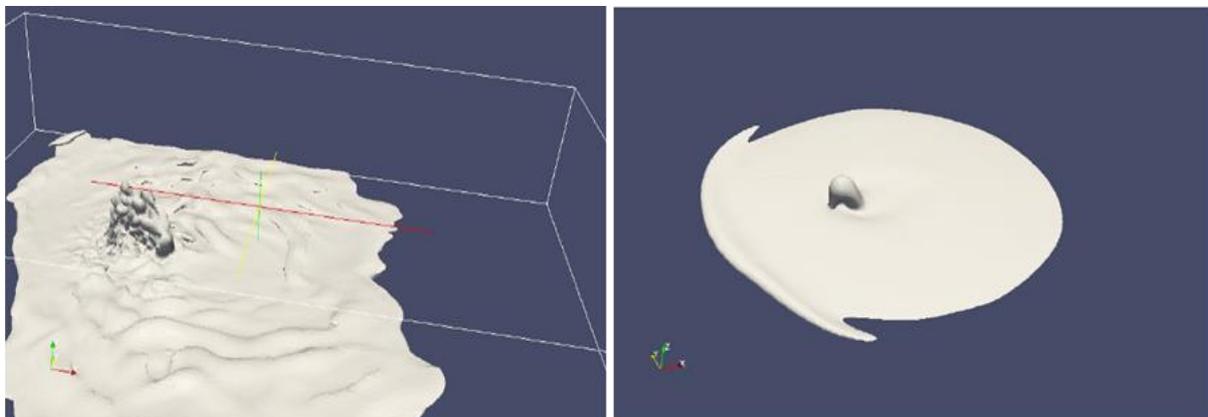
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523 **Overall behaviour of the CO₂ cloud**

524 Figures 7 (a) and (b), show the contours of 1% CO₂ concentration at different times after the dense gas
 525 cloud was released. The gravity slumping effect on the dispersion can be seen clearly in both figures,
 526 particularly at the earlier times where the flow vectors within the highly concentrated gas cloud
 527 exhibit a strong downward bulk motion. Initially the buoyancy generated forces and pressure
 528 gradients arising from density differences between the CO₂ gas cloud and its environment lead to a
 529 bulk motion that causes the dense cloud to spread in all directions near the release location (including
 530 a lateral spreading, as well as an upwind spreading against the prevailing wind direction).
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533 Figure 7. Overall view of mean concentration contours $\beta = 1\%$ at different times (a) earlier (71 s) and
 534 (b) later times (130 s).
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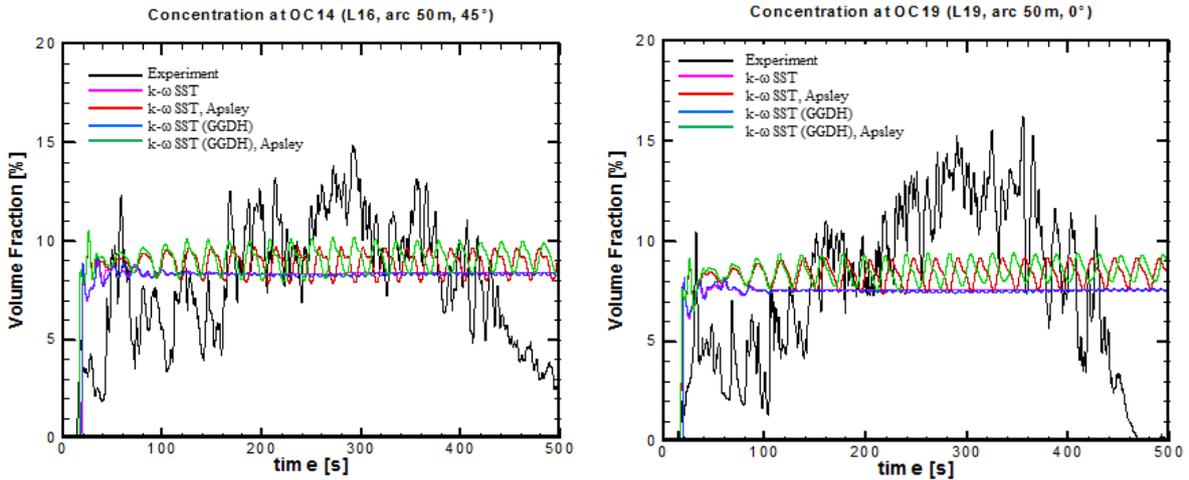
536 Further downstream from the release, the bulk flow resulting from the gravitational slumping weakens
537 as diffusion, mixing, and entrainment between the CO₂ gas cloud and the ambient air reduces the
538 negative buoyancy effects of the cloud and strengthens the influence of the externally imposed
539 velocity field on the transport and dispersion of the cloud. Whereas CO₂ was ‘sprayed’ upwards from
540 the jet nozzle at an initial velocity and subsequently became rapidly diluted with the surrounding air,
541 the initial velocity was still predominant resulting in the main tendency of CO₂ to move upwards, the
542 ‘spray’ has a noticeable horizontal component of around 2 m/s in magnitude (i.e. the initial release is
543 inclined to the crater wall at an angle of $\theta=70^{\circ}$). On top of the jet being inclined, the motion is
544 expected to be affected by the wind coming across, however in this case, the wind being relatively of
545 lower velocity, its motion was only slightly affected by it. The second phase of gravity slumping
546 which marks heavy gas dispersion is most significantly different from that of a passive gas. Due to
547 gravity, this phase was characterized by the slumping or collapse of the cloud towards the ground
548 level. It continues to entrain ambient air, which resulted in a reduced density, i.e. a reduced density
549 deficit. The loss of momentum of the jet is in a sense much faster than its dilution, creating the
550 prevalence of gravity effects in the movement of the cloud. Subsequently, the CO₂ cloud dispersion
551 goes into a third phase of gravity current, where the cloud pushes radially outwards, at a relatively
552 steady speed. As the cloud volume grows wider and larger, it is further diluted by the ambient air. The
553 decrease in concentration is gradual and the height of the CO₂ cloud decreased slightly. Finally, as the
554 cloud density decreased, the negative buoyancy effects disappear and the CO₂ plume motion
555 resembles that of passive gas.

556

557 **Comparison between the measured and predicted CO₂ concentrations by different variations of** 558 **the k- ω SST models and wall functions**

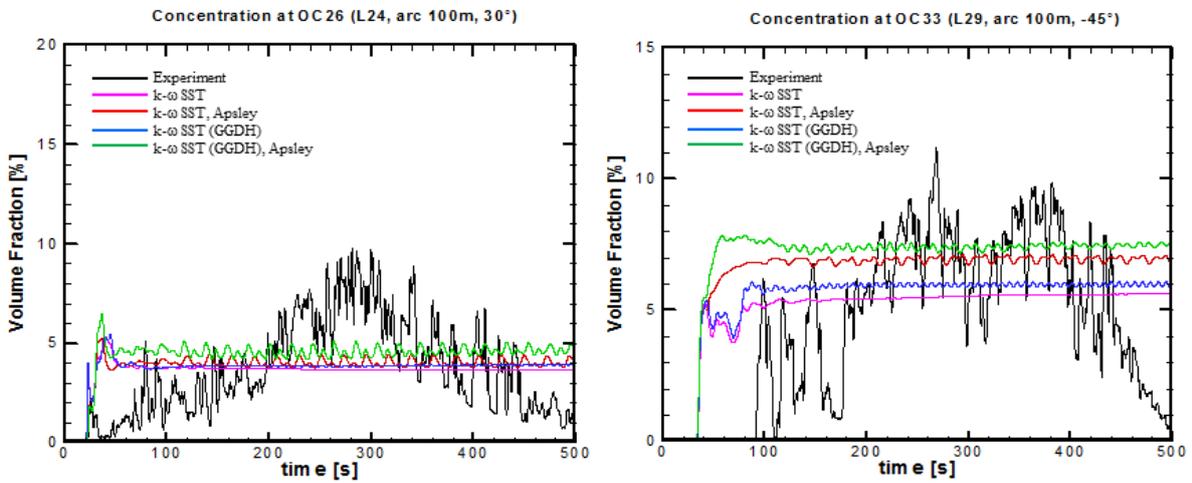
559 To further evaluate the potential improvement in the predictions which can be achieved by the use of
560 GGDH and improved wall functions extended to arbitrary wall roughness by Apsley (2007),
561 predictions have been carried out for the Case Study 3-base case for the k- ω SST with GGDH, k- ω
562 SST with the Apsley wall function and k- ω SST with both the GGDH and Apsley wall function. The
563 predictions are shown in Figures 8 to 11, along with the experimental measurements. When
564 examining the comparison with the measurements, one should bear in mind that the predictions are
565 conducted with the pseudo source which approximate the outflow conditions when the release has
566 reached quasi steady state. Hence one should not expect the predictions to capture the transient
567 characteristics of the actual dispersion. However, the numerical simulation predicted some
568 fluctuations in the concentration of CO₂ in the near fields, which can be attributed to the interplay
569 between the jet momentum, buoyancy and the gravity. These fluctuations are almost negligible in the
570 far field. Overall, the comparison indicates that although for some monitoring points the predictions
571 by the k- ω SST with both the GGDH and Apsley wall function are in closer agreement with the base
572 k- ω SST model, in the majority of the cases, the base k- ω SST model actually delivered predictions
573 that more closely match the experimental data. The comparison supports our choice in using the base
574 k- ω SST model for all the predictions and wall functions based on the roughness of the ground in the
575 present study following Richards (1993).

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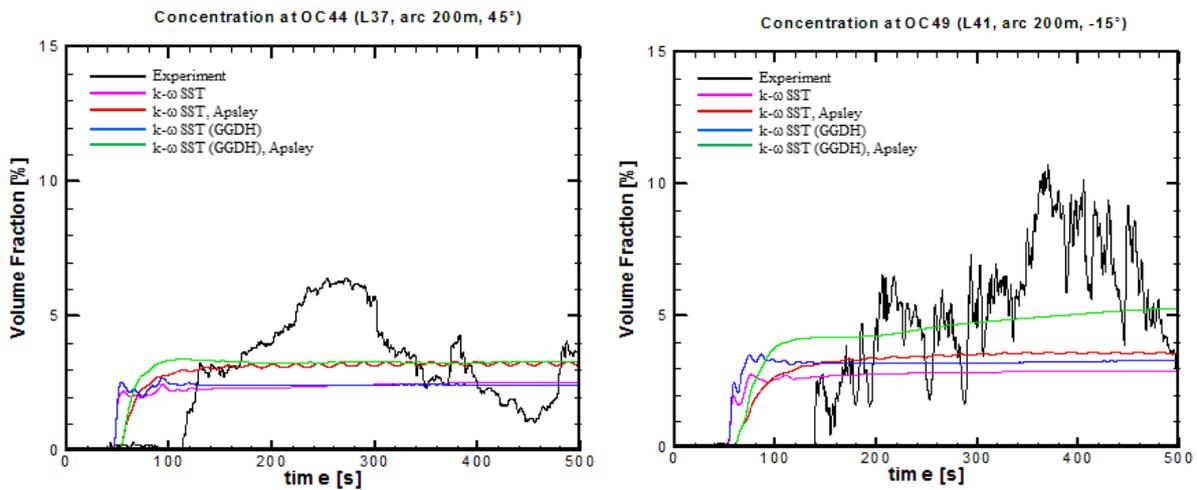
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Figure 8. Comparison of the measured and predicted concentrations with different model combinations on the arc 50 m.



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Figure 9. Comparison of the measured and predicted concentrations with different model combinations on the arc 100 m.



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Figure 10. Comparison of the measured and predicted concentrations with different model combinations on the arc 200 m.

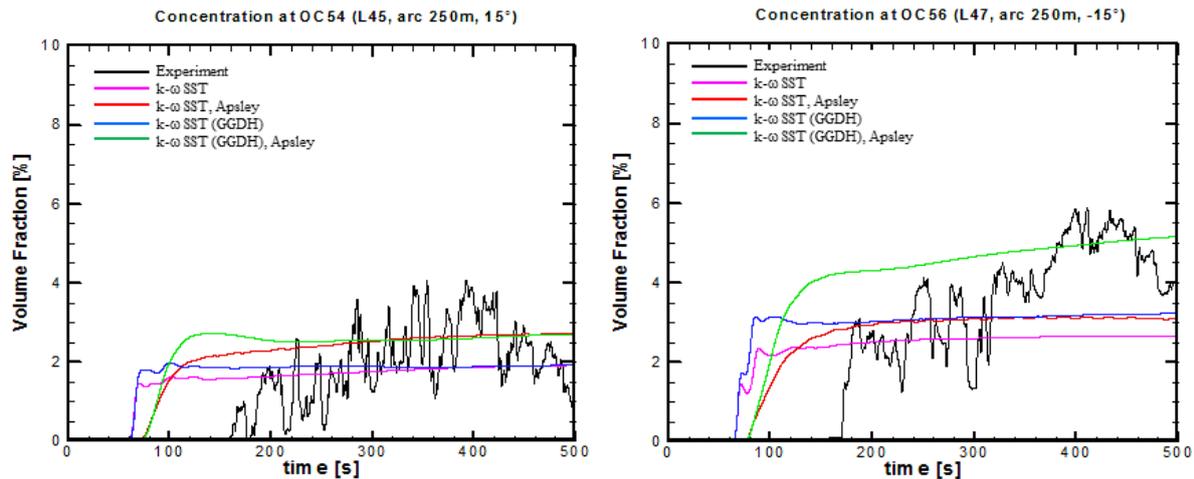
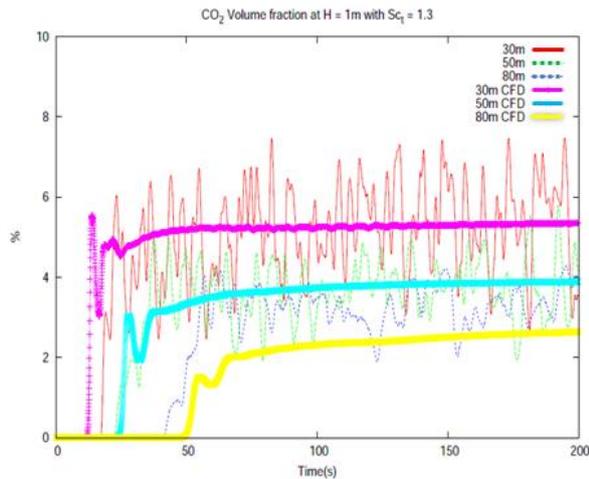


Figure 11. Comparison of the measured and predicted concentrations with different model combinations on the arc 250 m.

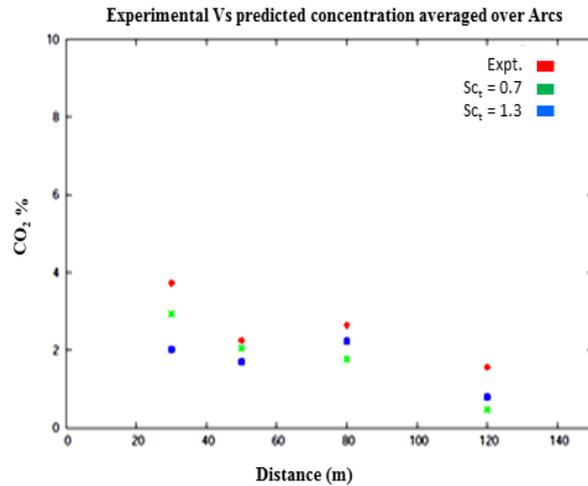
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Further comparison between the measured and predicted CO₂ concentrations

596 As previously stated, the CO₂ concentration levels are of important concern in the safety context. A
 597 minimum value of interest could be as low as 1%, which can make some people feel drowsy.
 598 Generally, numerical predictions of concentration are seen to be in quite good agreement with the
 599 experimental data. It is noted that in interpreting the level of agreement between the predictions and
 600 the measurements, one needs to recognize that the measurements of the short time-averaged
 601 concentrations obtained correspond to one realization of the instantaneous dispersion, whereas the
 602 model predictions correspond to an ensemble-averaged concentration (and as such, is associated with
 603 an average over an ensemble of realizations of the instantaneous dispersion). Because the model
 604 prediction is compared to one realization out of the ensemble of all possible realizations, the
 605 discrepancy between the model predictions and the experimental measurements cannot be smaller
 606 than the expected fluctuations of the individual realizations about the ensemble mean. As a
 607 consequence, it is impossible for even a perfect model to give a greater precision in its predictions of
 608 the concentration than the expected concentration variability from realization to realization in the
 609 atmospheric dispersion. As shown in Figures 12 and 13, both the predicted CO₂ volume fraction and
 610 the predicted and measured CO₂ concentration averaged over arcs by the RANS approach are in
 611 reasonably good agreement with the measurements. As noted by Clever et al. (2007) for the
 612 simulations of large scale dispersions of a liquefied natural gas spill, there appeared to be a degree of
 613 consensus that the better of the more, practical models (box or similarity models) should be within a
 614 factor of 2 of the observed concentrations for a straight-forward situation within the bounds covered
 615 by the experimental data. Similar comments were also reported by Hanna et al. (1993) and Daish et al.
 616 (2000). The agreement between the present predictions and the data averaged over the arcs are indeed
 617 within a factor of 0.5.
 618



619
620
621 Figure 12. Comparison between the predicted
622 and measured CO₂ volume fraction.
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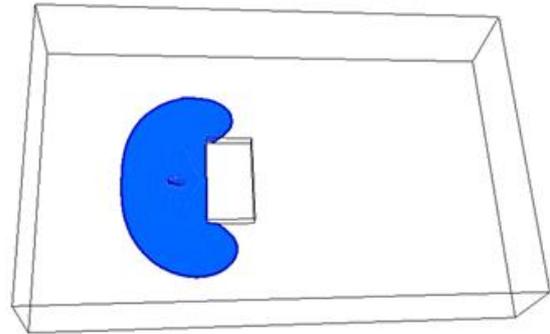
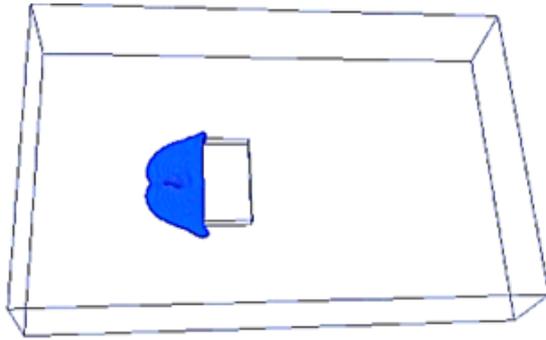
624 Figure 13. Comparison between the predicted and
625 measured CO₂ concentration averaged over arcs.
626

624 Obstacle effects

625 A series of studies have also been conducted to investigate the effects of obstacles on CO₂ dispersion.
626 While a more detailed paper about these studies is under preparation, some snapshots of results
627 involving a commercial building using release conditions in Case Study 3 as the baseline release are
628 included here to illustrate the application of the code. The commercial building is a single storey large
629 building (representing a factory or warehouse) of dimensions 60 m wide by 30 m deep by 8 m high
630 and placed with its front face 25 m downwind of the release. For simplicity, all windows or doors are
631 assumed to be closed. The simulations were conducted for a wind speed of 1 m/s at a 10 m reference
632 height. The simulations were conducted in two steps: the computations were run at first before the
633 CO₂ was released to reach a steady-state atmospheric flow field. The flow field thus obtained was
634 then used as the initial conditions to study the dispersion of CO₂ following the release.
635

636 The 2% iso-contours of CO₂ concentrations at t = 30 s and 300 s are shown in Figures 14 and 15
637 respectively. These overall views of the cloud show that the cloud wraps around the obstacle but not
638 over it. In other words, high concentrations (1.5% and over) are contained or blocked by the
639 commercial building which limits the spreading downstream of the obstacle but increases the spread
640 laterally and upwind of the release point against the prevailing wind direction.
641

642 The commercial building also has considerable impact on the wind flow. Prior to the release, the
643 buildings create a considerable velocity deficit in the atmospheric boundary layer, with recirculation
644 zones lee- and wind-ward as well as to the side of the obstacle, the wind profile only returns to an
645 equilibrium more than 50 m downwind. The recirculation zones are clearly illustrated by the
646 streamlines shown in Figure 16. The interaction of the cloud with these recirculation patterns is most
647 distinct in the decrease of the recirculation intensity and creation of the low velocity “calm” zones
648 where the CO₂ can spread slowly (observe the flattened velocity profiles around the buildings).
649

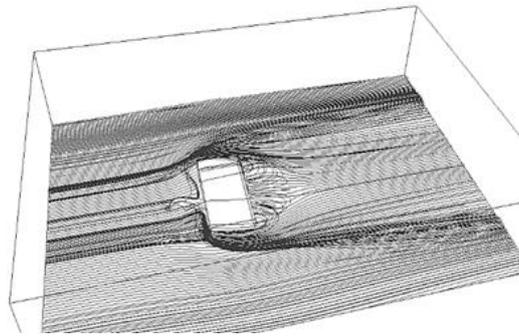


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Figure 14. The predicted 2% CO₂ iso-contour at t = 30 s

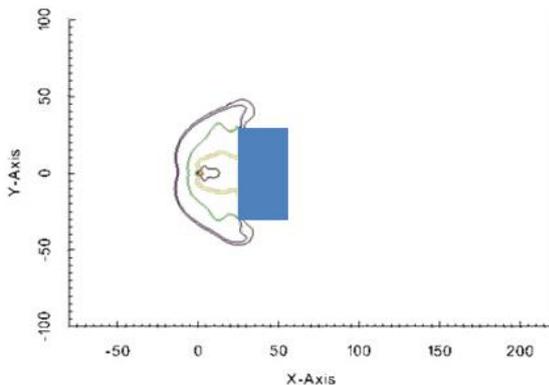
Figure 15. The predicted 2% CO₂ iso-contour at t = 300 s



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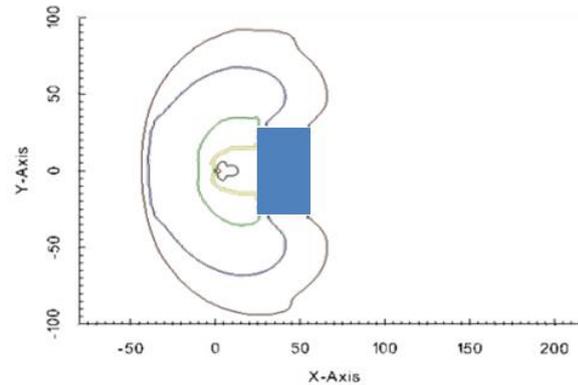
Figure 16. Streamlines around commercial building.

658 Top-views in Figures 17 and 18, illustrate the horizontal reach of the CO₂ cloud at t = 30 s and t = 5 min, respectively. Iso-concentration contours of 20%, 10% and 7% are observed in the immediate vicinity of the release and form a circular spreading pattern away from the obstacle. However, the building blocks the CO₂ cloud and dilute it at the building edges as the cloud tries to wrap around it. 661 A cross-wind cut plane view, right after the obstacle (2 m downwind from the obstacle), shown in figure 19 showing only concentrations levels of 0.5% CO₂. A perpendicular view to the former is shown in figure 20. Both help to illustrate the combined blocking/diluting effect of the obstacle. The 665 0.5% iso-contours only exist to the sides of the building further, and higher concentrations are clearly contained upstream of the building. 666



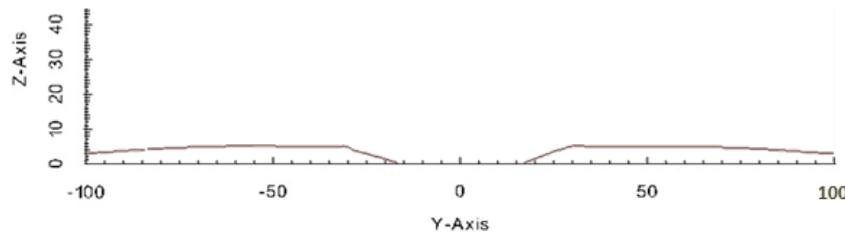
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668 Figure 17. Horizontal CO₂ concentration profile at
669 1 m height and 30 s. Red: 20 %, black: 10 %, yellow: 7%, green: 4%, blue: 1.5% and brown: 0.5%.

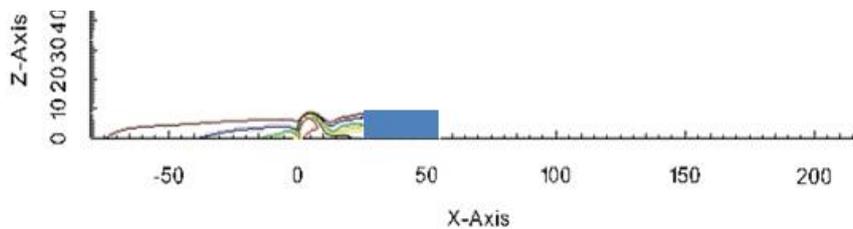


670 Figure 18. Horizontal CO₂ concentration profile at 1 m height at t = 5 min. Red: 20 %, black: 10 %, yellow: 7%, green: 4%, blue: 1.5% and brown: 0.5%.

671



672 Figure 19. Vertical volume fraction contours at 37 m downwind of the release point at $t = 5$ min.
 673 brown: 0.5%.
 674
 675



676 Figure 20. CO₂ iso-contour in the plane $x0z$ at $t = 300$ s. red (inner most): 20%, black: 10%, yellow:
 677 7%, green: 4%, blue: 1.5% and brown: 0.5% (outer most) .
 678
 679

680 Conclusions

681 CO₂FOAM, a dedicated solver for CO₂ dispersion previously developed within the frame of the open
 682 source CFD code OpenFOAM® by the authors' group has been further extended to include the
 683 homogeneous relaxation model which is more suited to account for the presence of solid CO₂ within
 684 the release. The code offers both RANS and LES approaches, but only the RANS approach is
 685 reported in the present paper. In order to better capture the interaction between the dispersed CO₂ and
 686 the ABL-specific boundary layer, a compressible form of the $k-\omega$ SST turbulence model is used in
 687 conjunction with ABL-specific wall-functions. Predictions have been conducted for Case Study 3 in
 688 the COOLTRANS research programme. The CO₂ was released through a puncture in a buried pipe.
 689 The numerical simulations used the near-field dispersion studies from Wareing et al. (2014) as input
 690 data for the far field dispersion simulation. Evaluations of the different variations of the $k-\omega$ SST
 691 model suggested that the baseline $k-\omega$ SST model delivered the predictions that more closely matched
 692 the experimental measurements for the tests cases considered here.
 693

694 The predictions have achieved reasonable agreement with the data on the basis of a “blind validation”,
 695 giving confidence to the capability of CO₂FOAM to be used for quantified risk assessment involving
 696 far field CO₂ dispersions; especially for situations where screen tests with semi-empirical or integral
 697 models suggest more detailed consequence analysis. To illustrate the application of the code, further
 698 predictions were conducted using the same release conditions as the baseline case but for scenarios
 699 involving a commercial building placed with its front face 25 m downwind of the release. The CO₂
 700 cloud wraps around and passes the building, generating turbulence as well as large flow oscillations
 701 and recirculation zones. The increased turbulence level due to the existence of the building would lead
 702 to more intense mixing and contribute to the dilution of the CO₂ cloud. This effect is, however,
 703 decreased by buoyancy forces as density stratification in the flow tends to suppress turbulence
 704 generation. The building has a blocking effect to the flow, where the cloud can be trapped temporarily
 705 or diverted towards particular zones.
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707 References

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